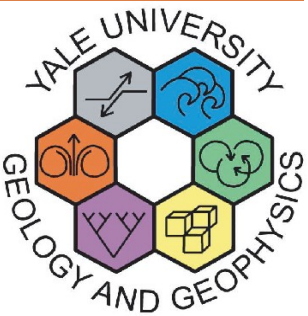


# Understanding Weak Low-Latitude SST Gradients and the Ocean Warm Pool Expansion in the Early Pliocene

Alexey Fedorov & Chris Brierley (Yale)

PP11G-08, AGU fall meeting 2009



Fedorov, A., C. Brierley, and K. Emanuel, (2010). Tropical cyclones, permanent El Niño and the climate of the early Pliocene. *Nature*, in press.

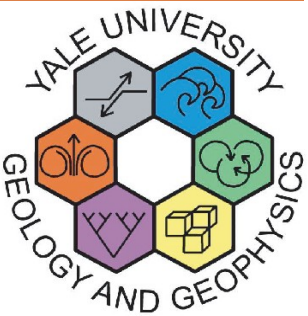
Brierley, C. and A. Fedorov, (2010). The relative importance of meridional and zonal SST gradients for the onset of the ice ages and Pliocene-Pleistocene climate evolution, *Under revision for Paleoceanography*.

Brierley, C., A. Fedorov, Z. Lui, T. Herbert, K. Lawrence and J. LaRiviere. (2009). Greatly expanded tropical warm pool and weakened Hadley circulation in the early Pliocene, *Science*, Vol. 323. no. 5922, pp. 1714 – 1718

Fedorov, A.V., P. Dekens, A. C. Ravelo, P. deMenocal, R. Pacanowski and S. G. Philander, (2006). The Pliocene paradox (mechanisms for a permanent El Niño). *Science* 312, 1437-1443.

## Alexey Fedorov & Chris Brierley (Yale)

Understanding Weak Low-Latitude SST Gradients and the Ocean Warm Pool Expansion in the Early Pliocene (PP11G-08, AGU fall meeting 2009)

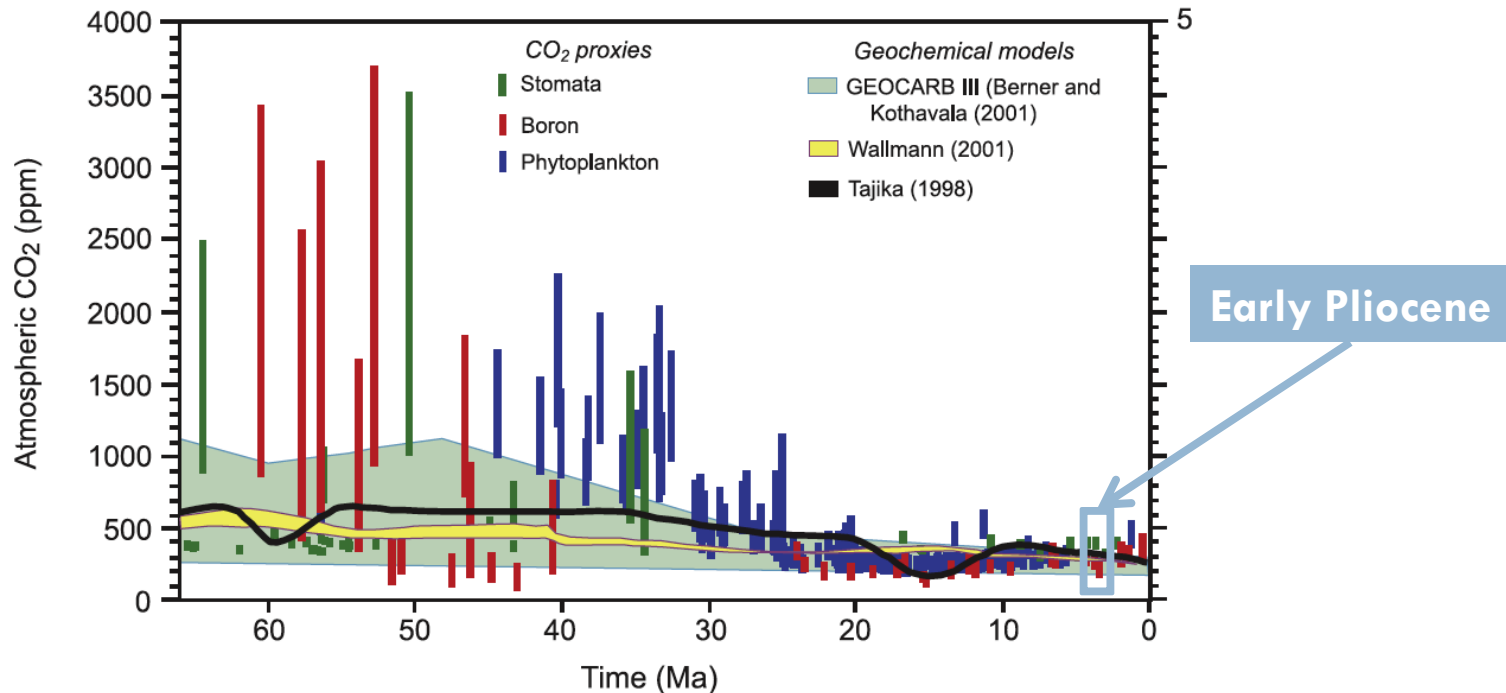


# Outline

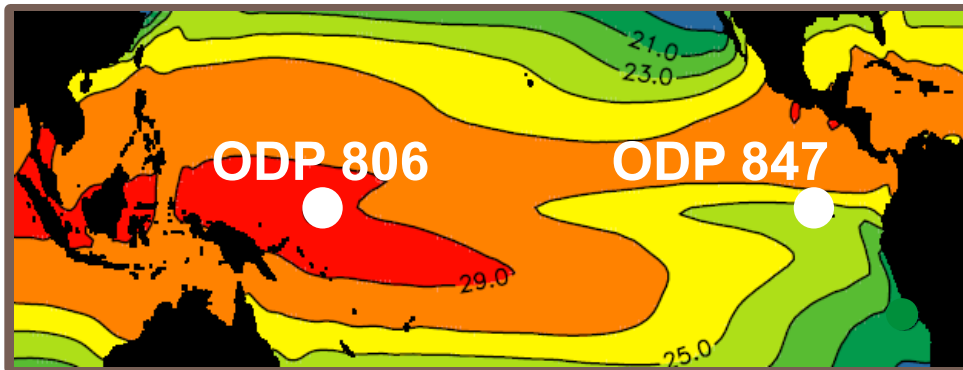
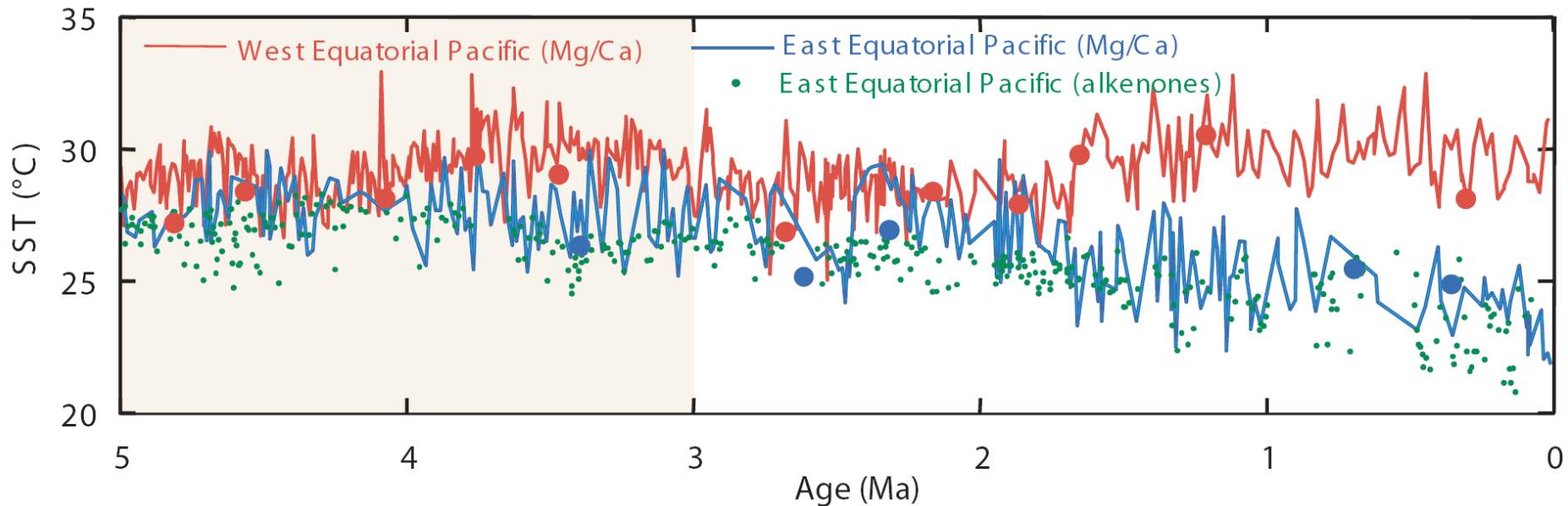
- Introduction to the early Pliocene climate
  - ▣ When & why should we care?
  - ▣ Tropical SST patterns
- Pliocene Paradox and missing heat transport?
- A tropical cyclone feedback?
  - ▣ The subtropical ocean circulation
  - ▣ Warming of the cold tongue
- Does this feedback explain the Pliocene warm pool?

# Why care about the early Pliocene?

- Natural global warming stabilization experiment
  - Pliocene CO<sub>2</sub> was 300 – 400 ppm
  - Continental configuration & orbital forcings similar

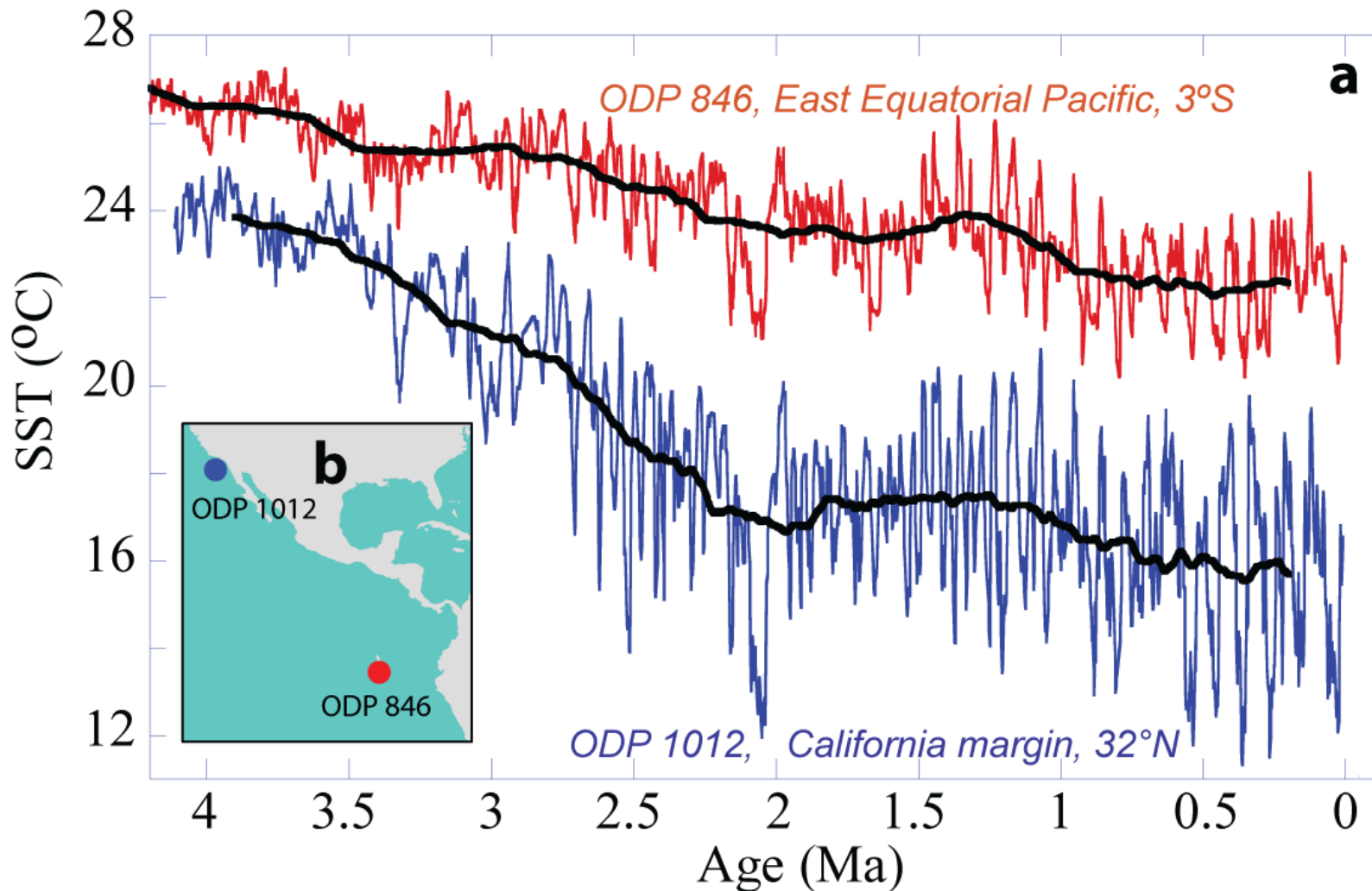


# Permanent El Niño (No Zonal SST Grad.)

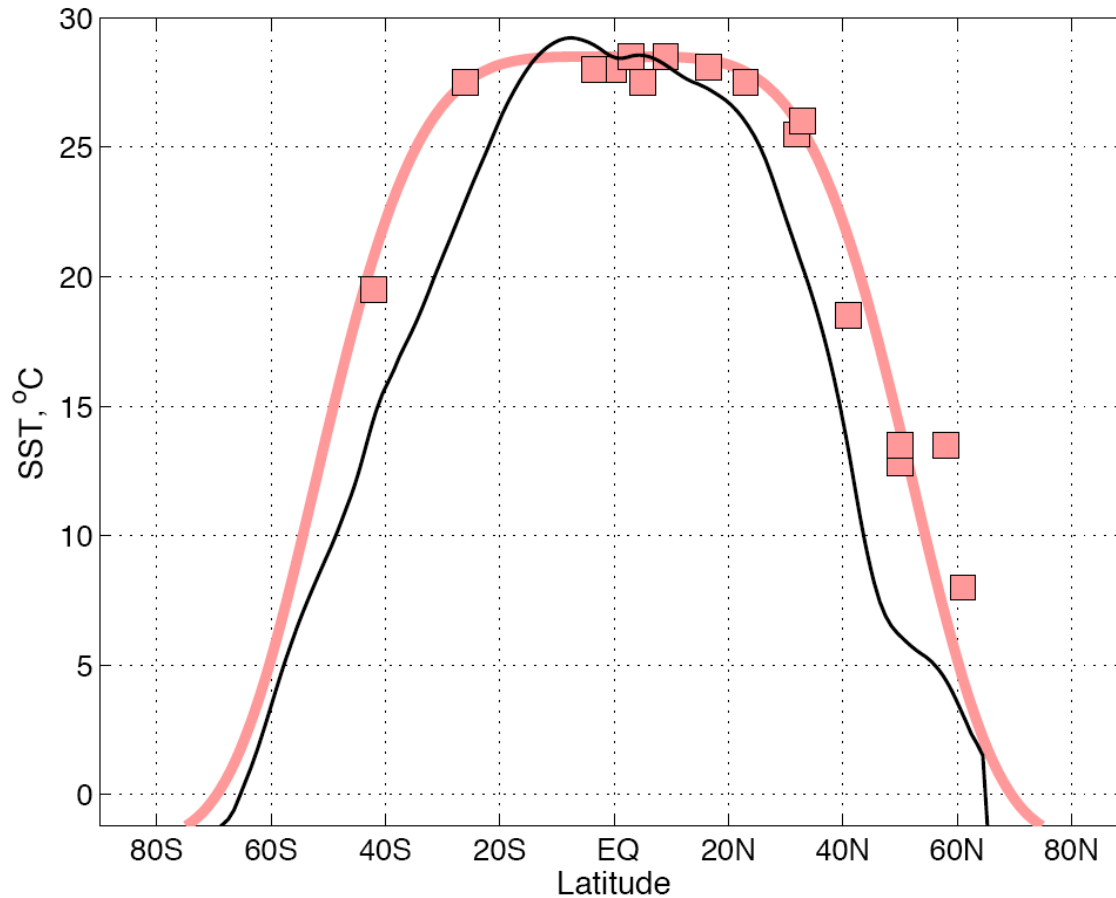


Fedorov *et al.* 2006. [The Pliocene Paradox \(Mechanisms for a permanent El Niño\)](#)

# Reduced Meridional SST Gradient



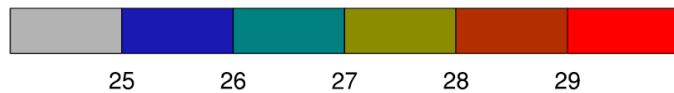
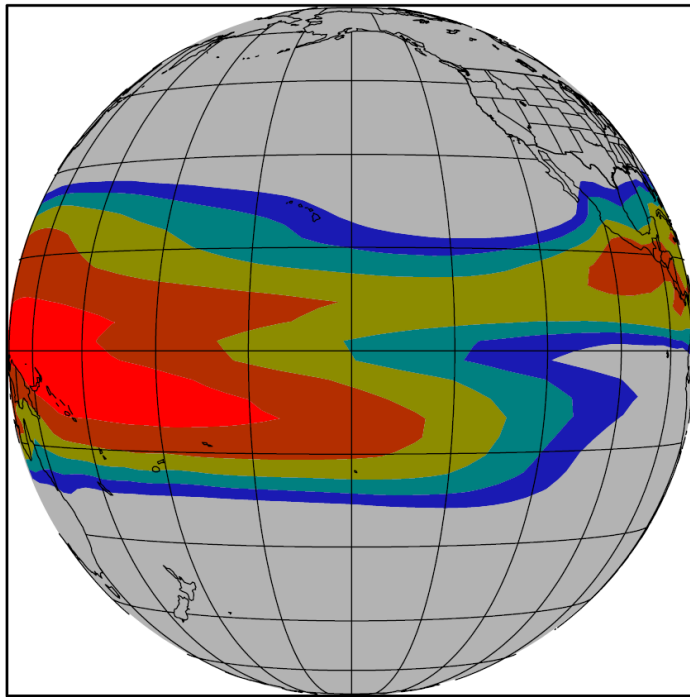
# Reconstructed Pacific SST profile



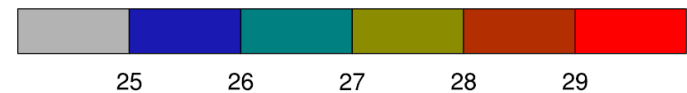
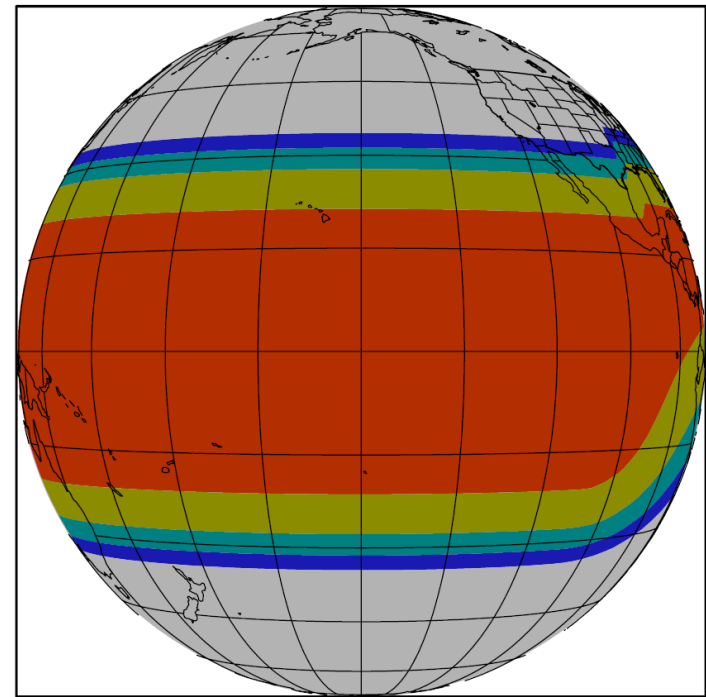
- Extend zonally across Pacific
- Shift meridionally to replicate the seasonal cycle

# Expansion of Warm Pool

(a) Present-Day SSTs



(b) Early Pliocene SSTs



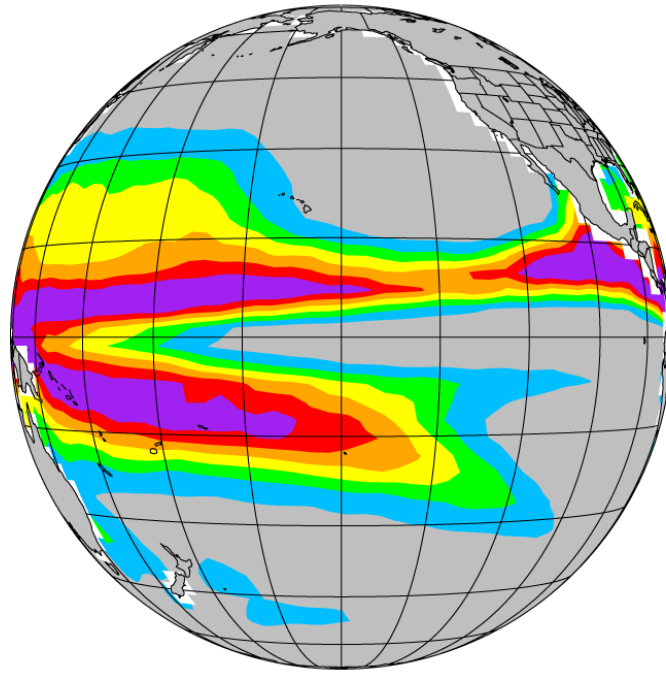


# Expansion of Convection

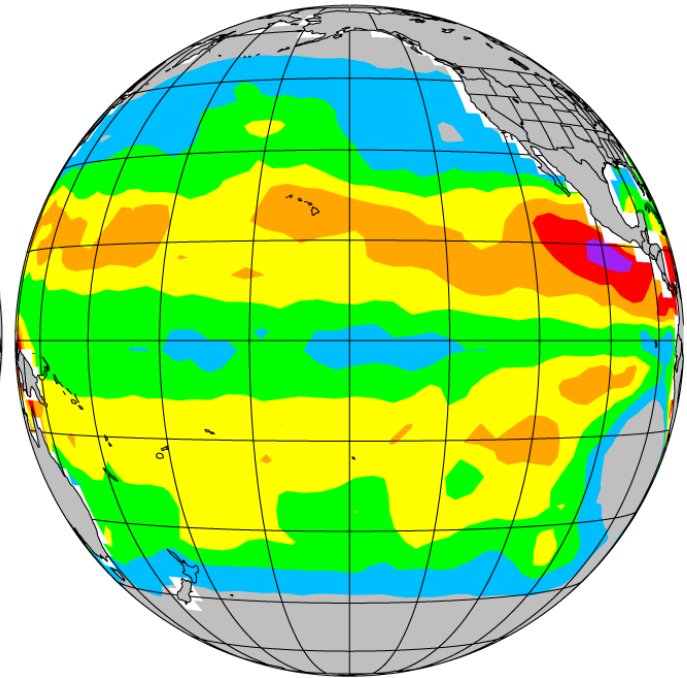
Use AGCM  
(CAM3 T85)  
with fixed  
SSTs to find  
impacts of  
warm pool

1990 CO<sub>2</sub>  
levels, modern  
orography,  
only SST (and  
sea ice) diff.

Model simulation of Present-Day

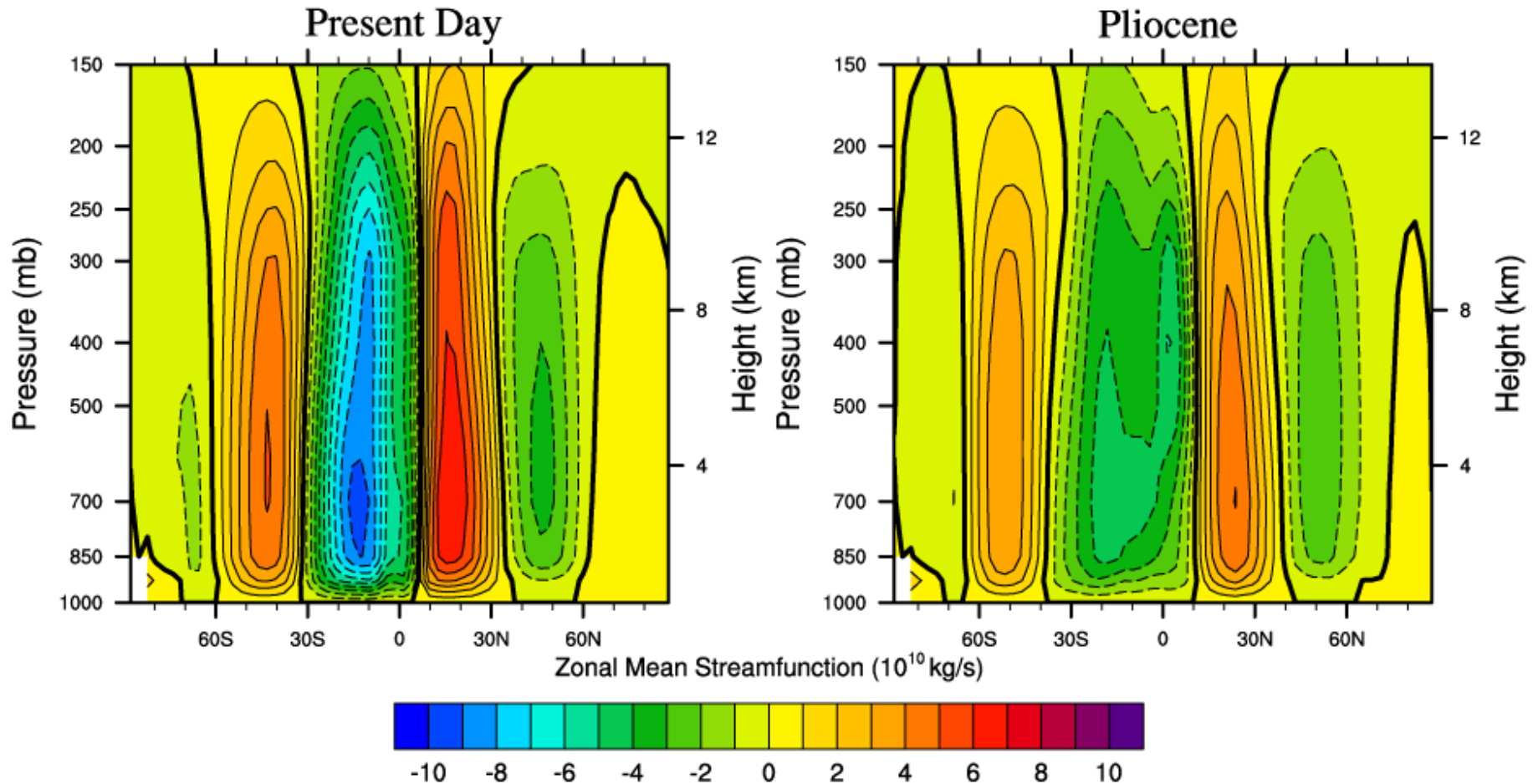


Model simulation of Pliocene



2 4 6 8 10 12  
Deep Convective Cloud Fraction (%)

# Weak Hadley Cell

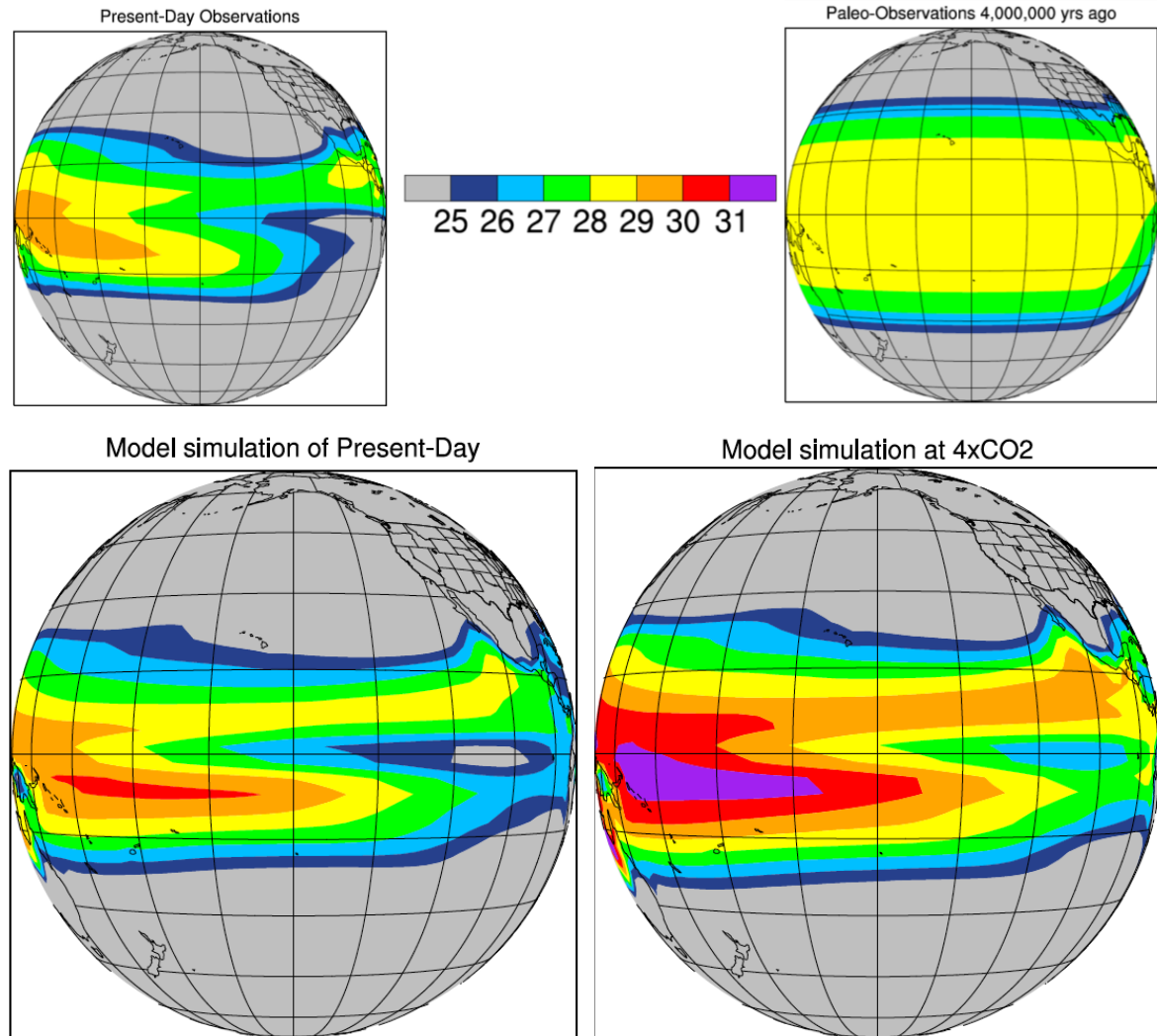


[Brierley et al., Science. 2009](#), [Fedorov et al., Science. 2006](#)

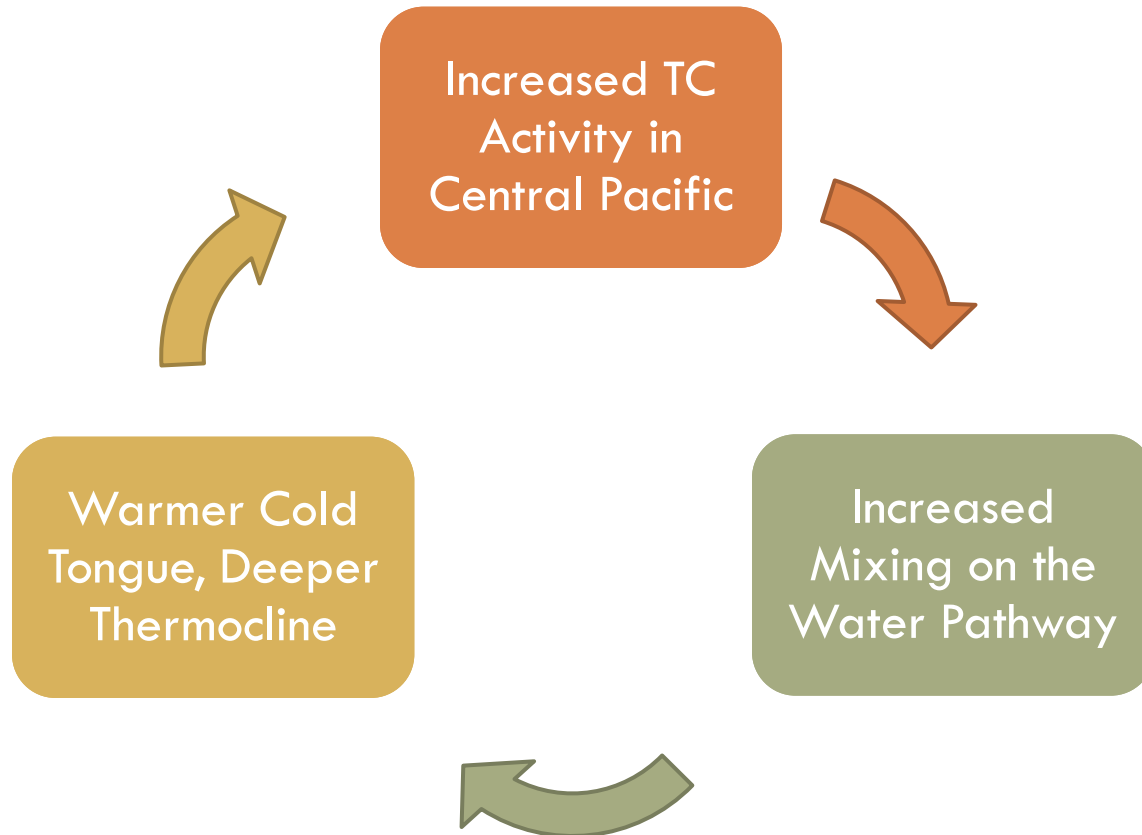
# The Pliocene Paradox

Pliocene warm pool unlike predicted response to added CO<sub>2</sub>

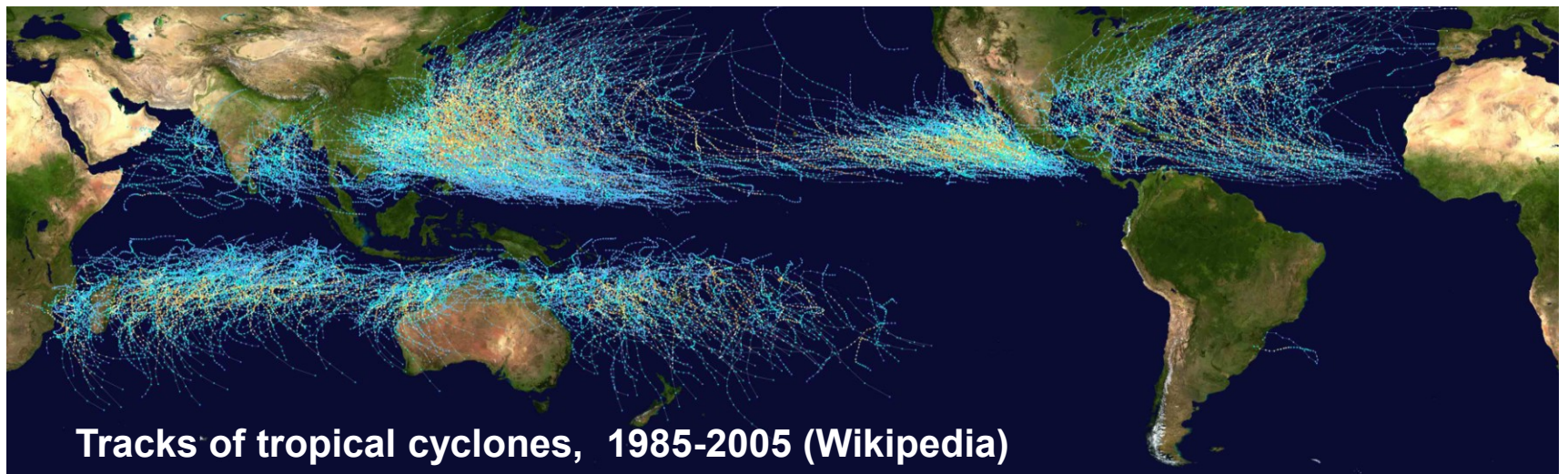
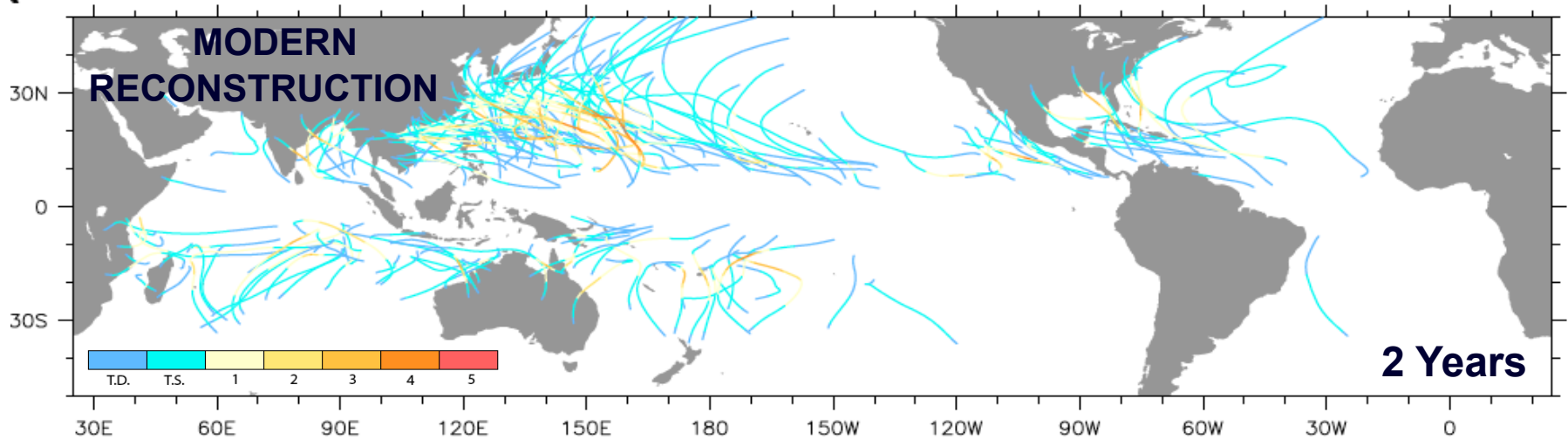
Maybe Earth has heat transport mechanisms not included in coupled GCMs?



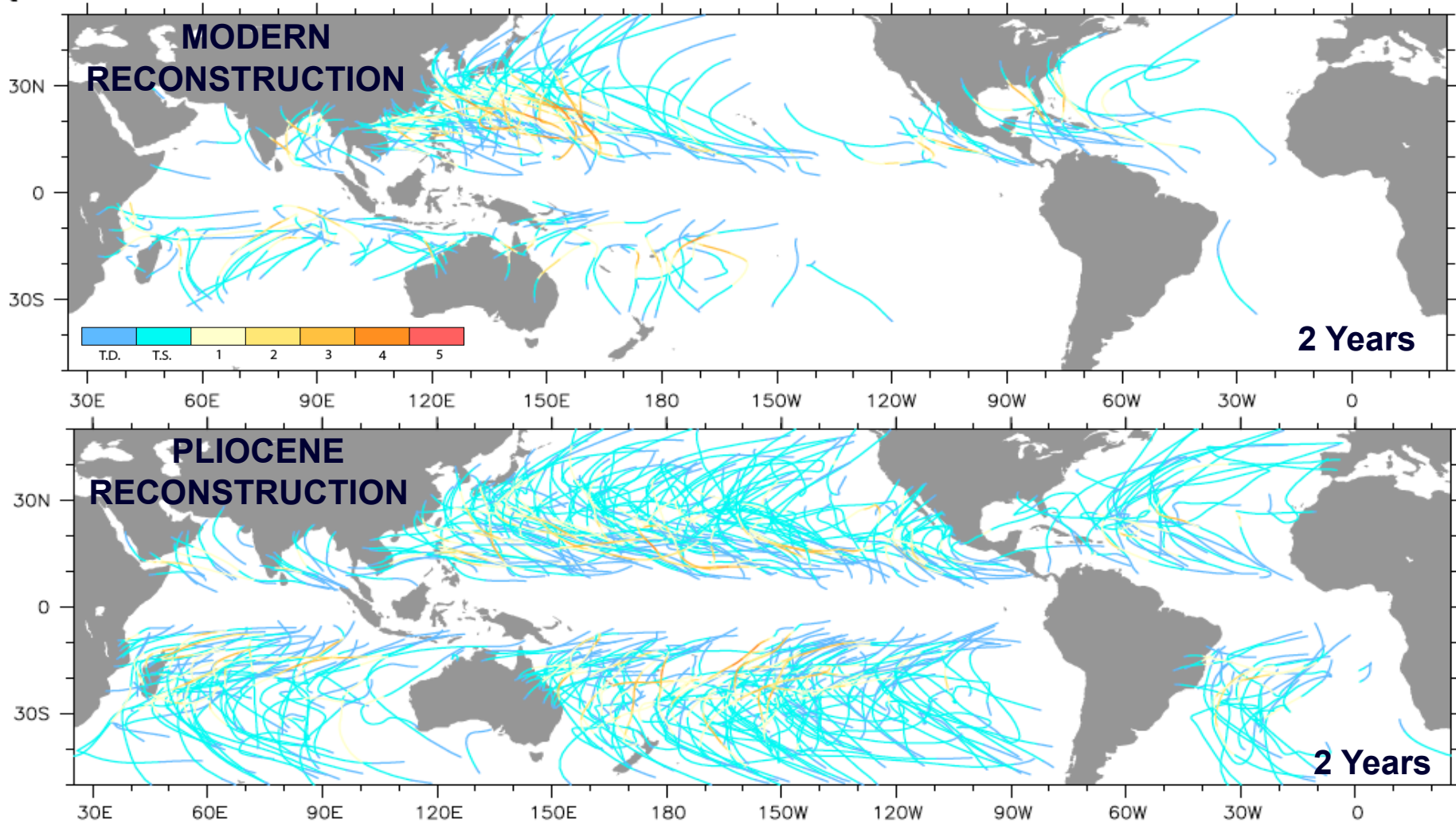
# Tropical Cyclone Feedback



# Synthetic Tracks for Present-day

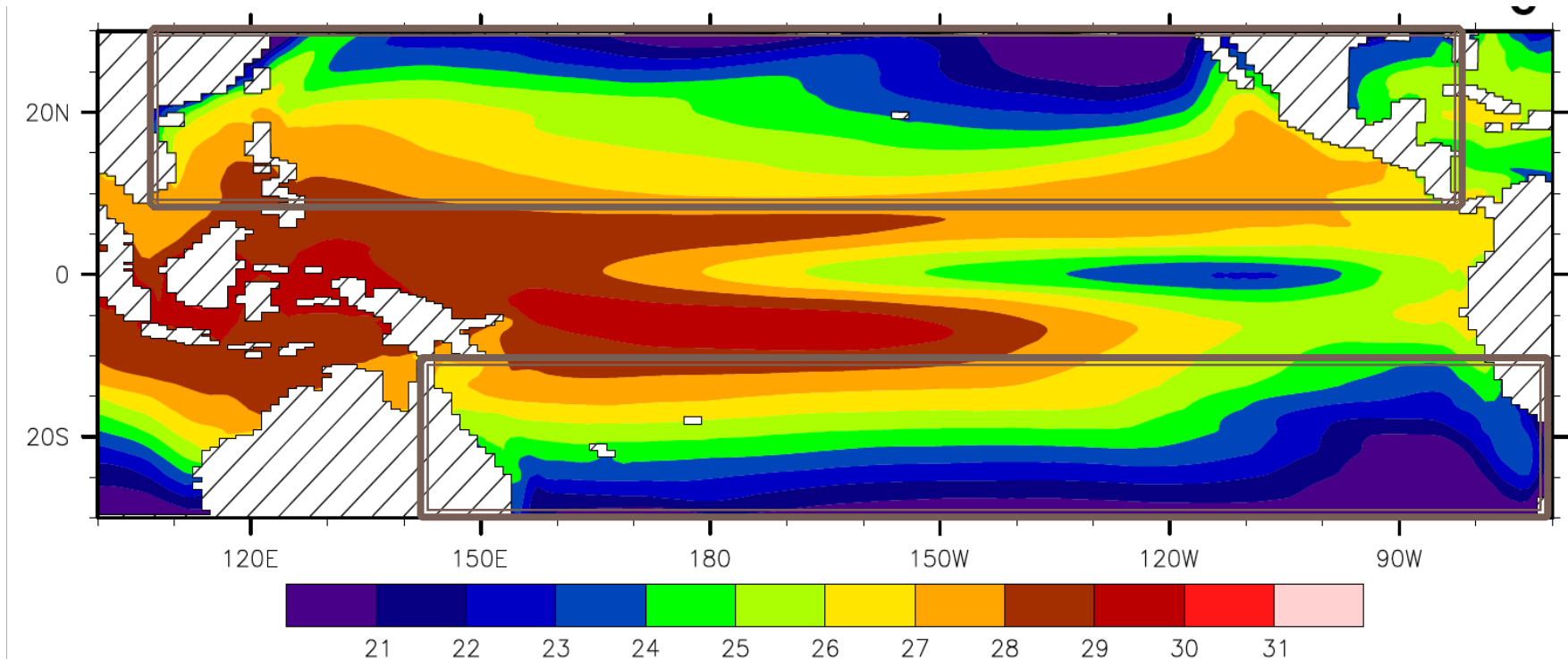


# Synthetic Tracks for Pliocene



# Modeling “tropical cyclone” mixing

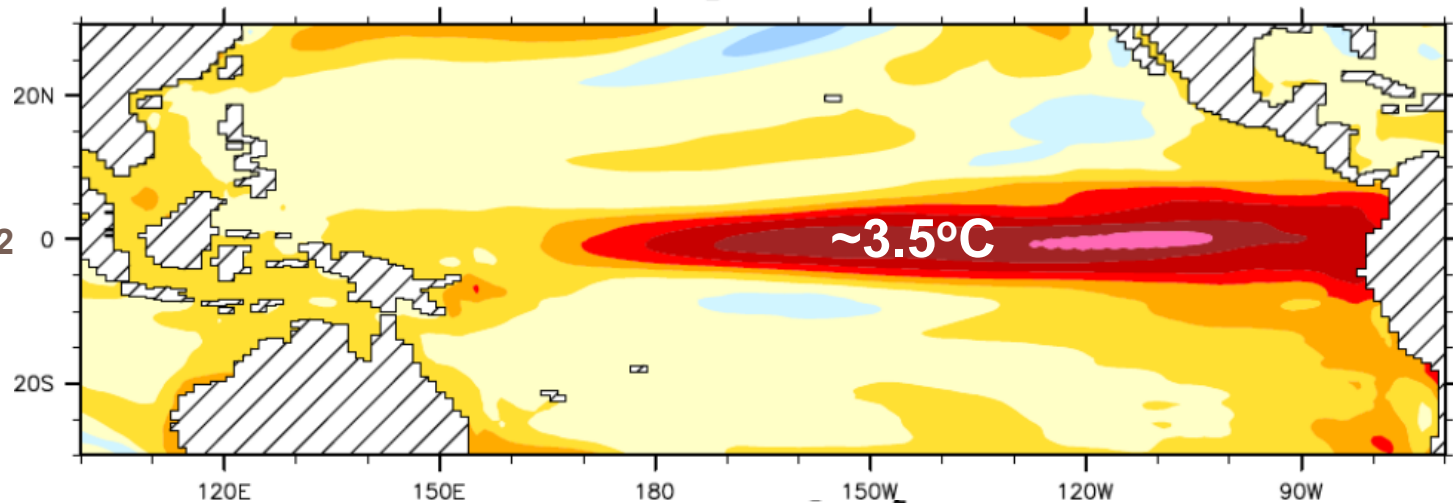
## Preindustrial SST



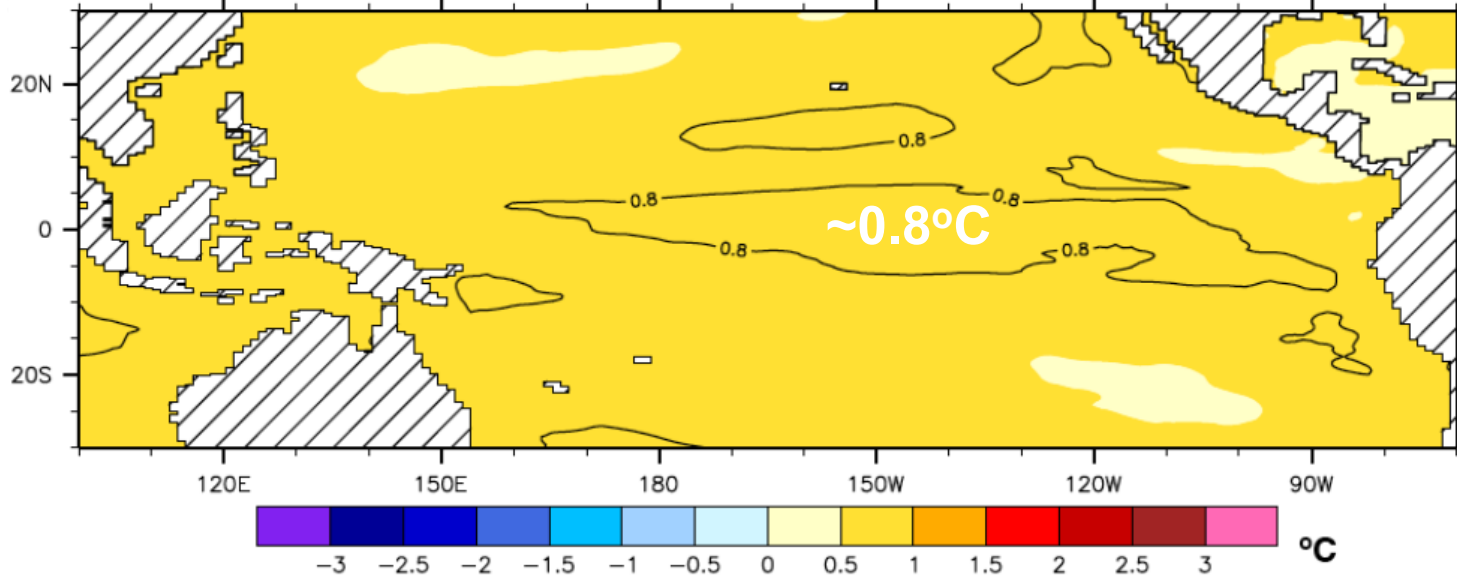
*We impose enhanced ocean mixing  
between 8-40° latitude ( $K_v=1\text{cm}^2/\text{s}$ )*

# Impact of “tropical cyclone” mixing

TC Mixing  
+ increasing CO<sub>2</sub>



Increasing  
CO<sub>2</sub> Only





# Conclusions

- The Tropical Pacific had a different SST distribution in the early Pliocene than at Present
  - ▣ One vast warm pool stretching from Indonesia towards California
- This vast warm pool created a sluggish atmospheric circulation.
- Sustaining the warm pool needs an additional physical process included in climate models
- Tropical cyclone feedbacks could be that process
- This feedback could be important in future projections