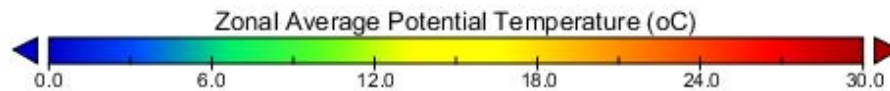
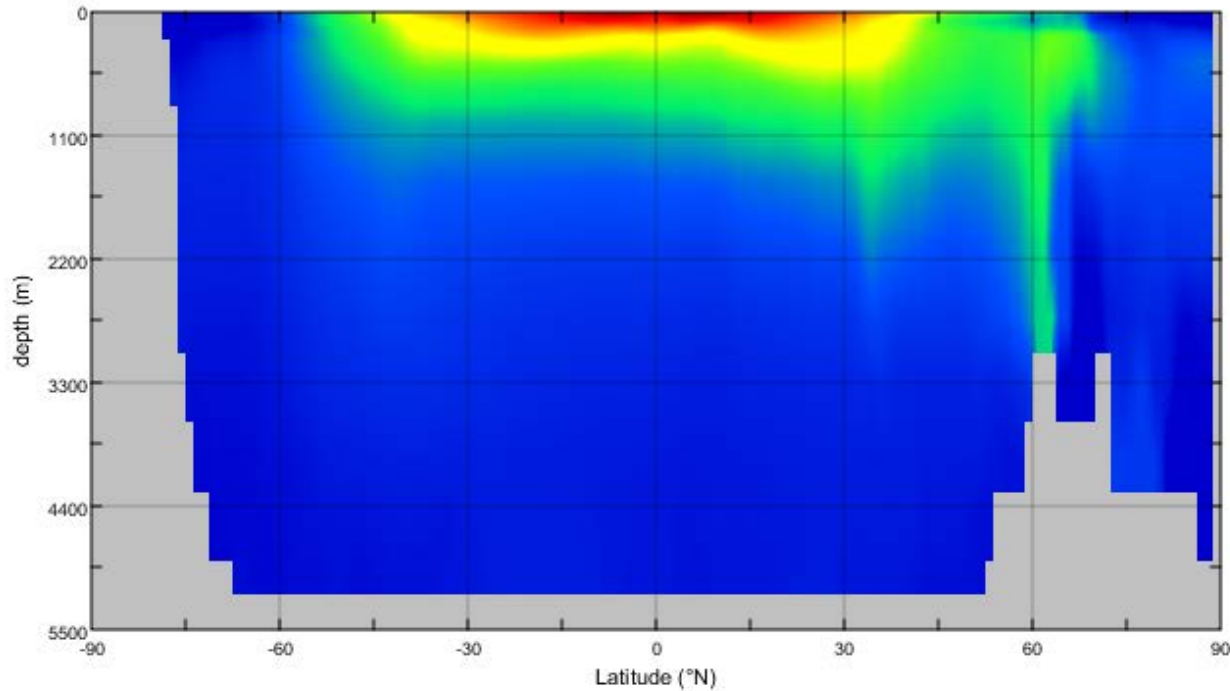


Vertical mixing and the ocean circulation

Chris Brierley,
Postdoc Luncheon, 14th April 2010

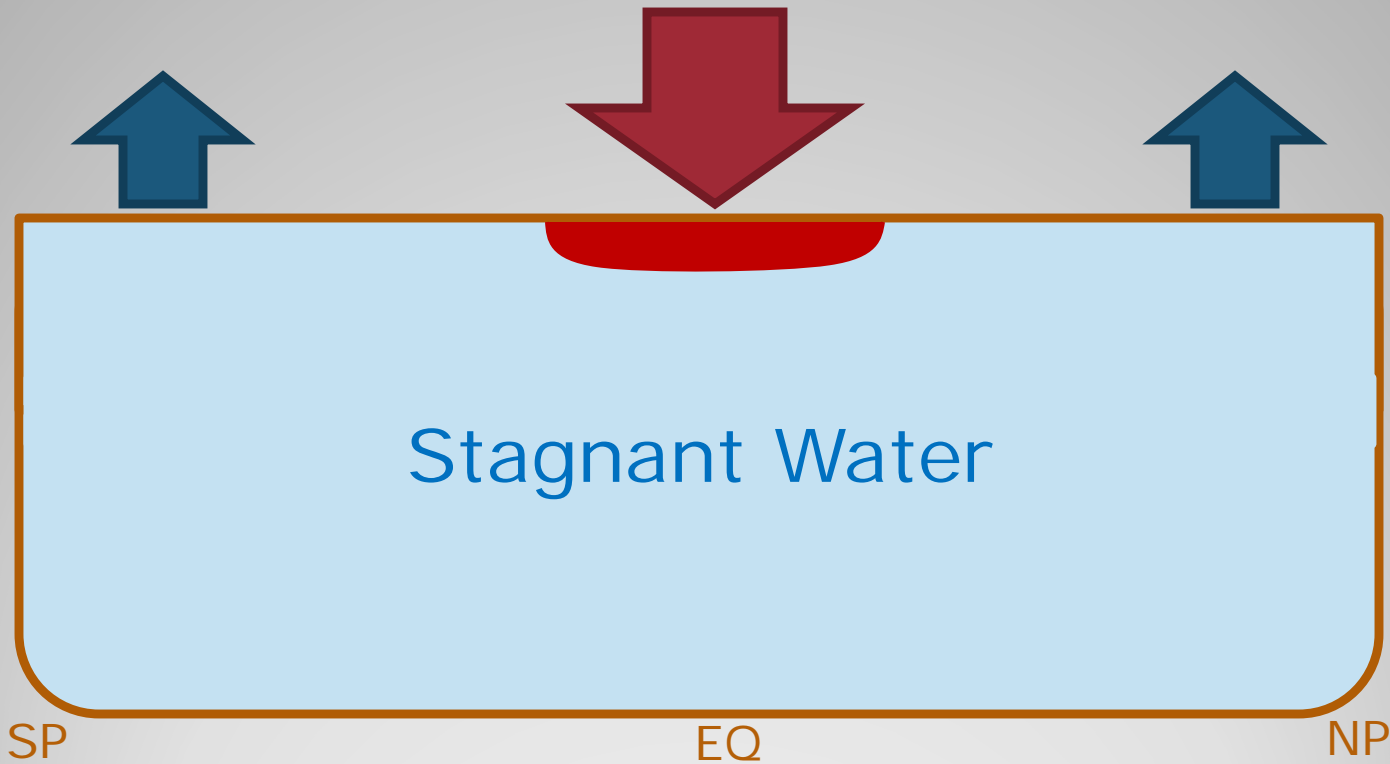
- Vertical mixing – what and why
- What is its global impact?
 - Mean climate
 - Climate change
- Changing sources of mixing
 - Hurricanes
 - Tidal dissipation

Outline



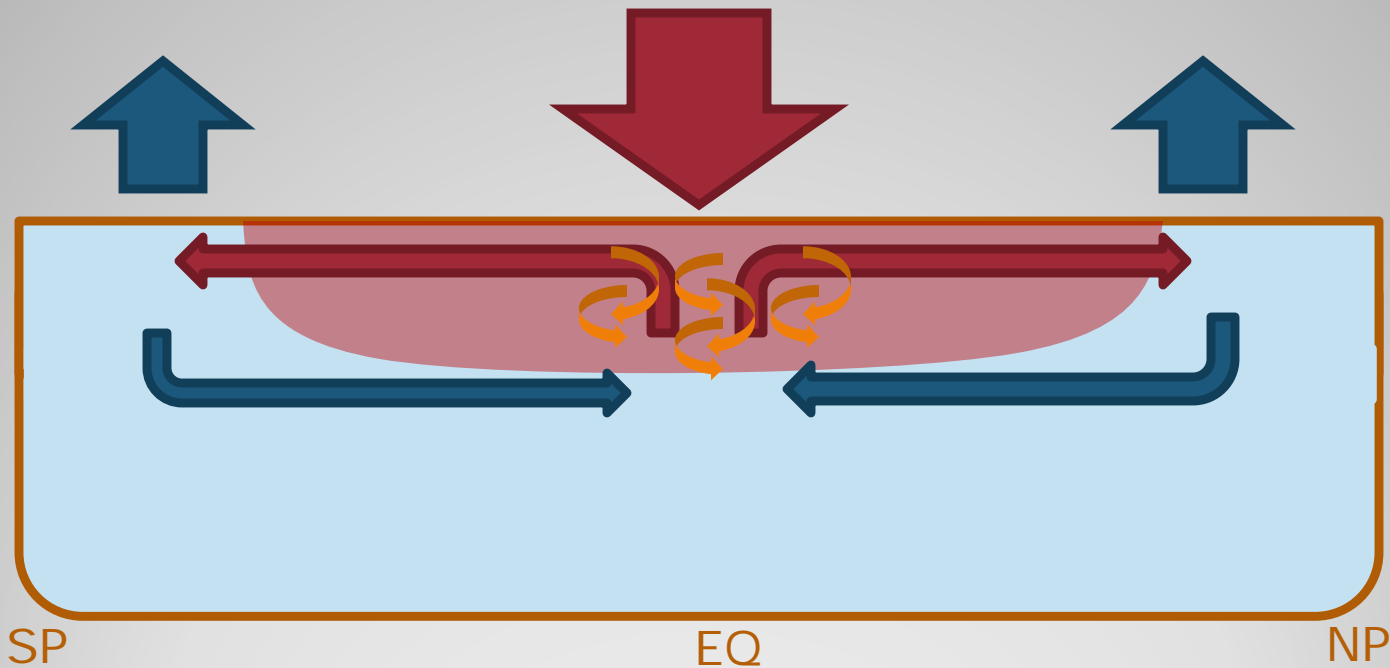
Zonal Mean Ocean Temperature

- *A circulation cannot be driven unless heat is input at a lower depth than it is lost at*



Sandström's Theorem

- Vertical mixing leads to smearing of water properties
- Effectively different vertical levels.

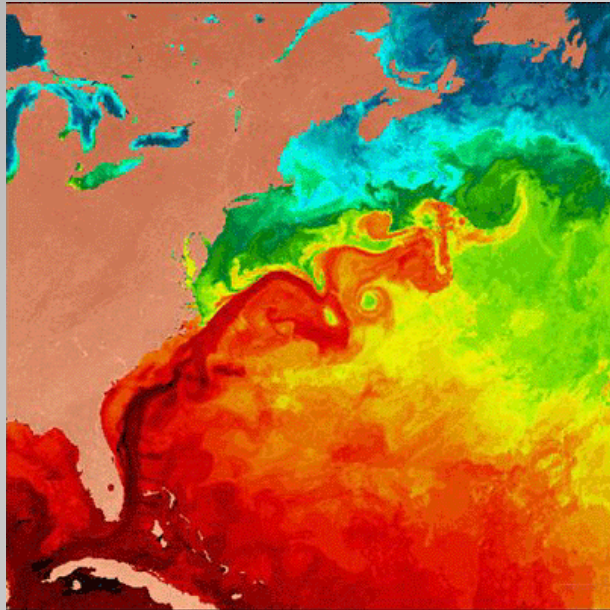


Mixing's impact

- What are the sources of mixing in the ocean?
 - Wind-driven stirring
 - Tidal generation of internal waves
 - *Biota*
 - *Hurricanes*

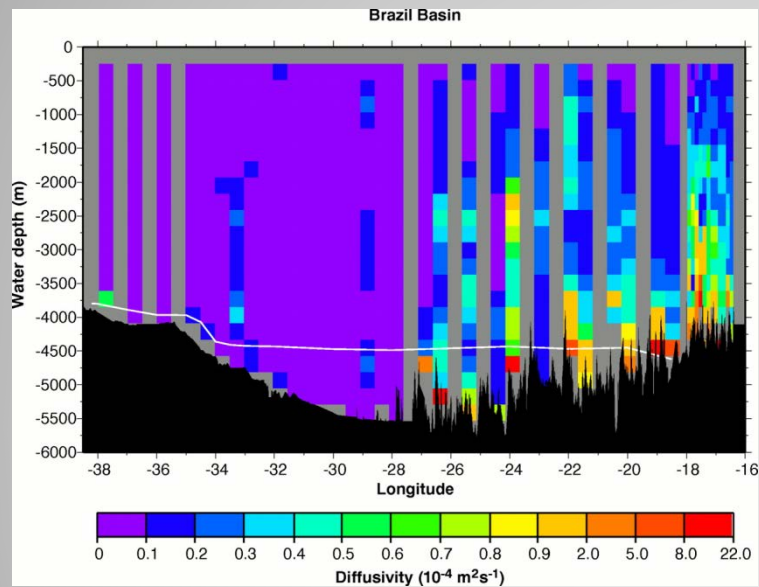
Sources of Ocean Mixing

- Wind-driven stirring
- Tidal generation of internal waves
- *Biota*
- *Hurricanes*



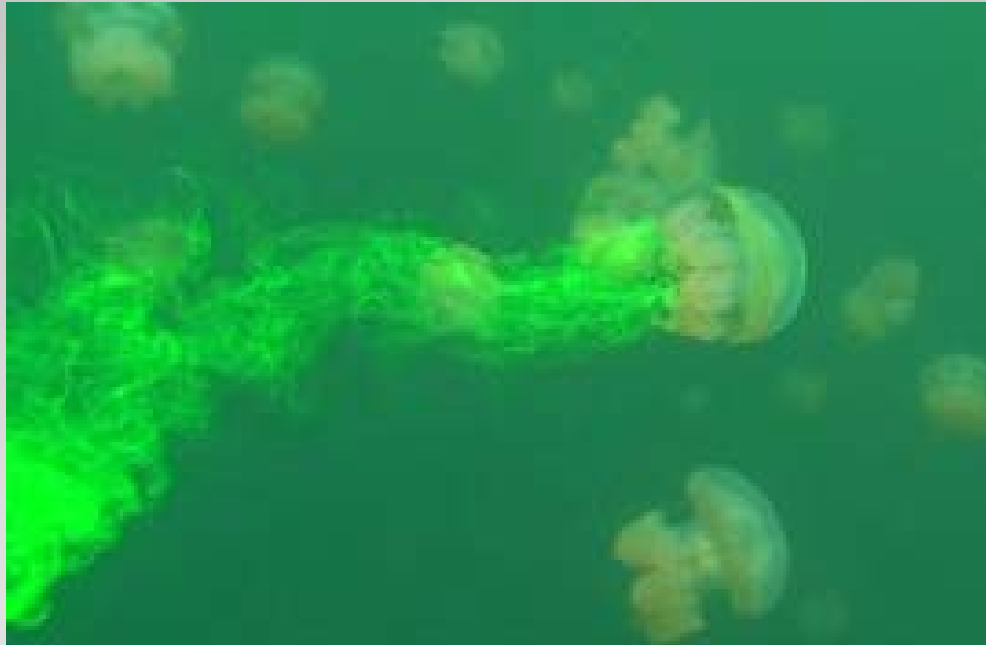
Sources of Ocean Mixing

- Wind-driven stirring
- Tidal generation of internal waves
- *Biota*
- *Hurricanes*



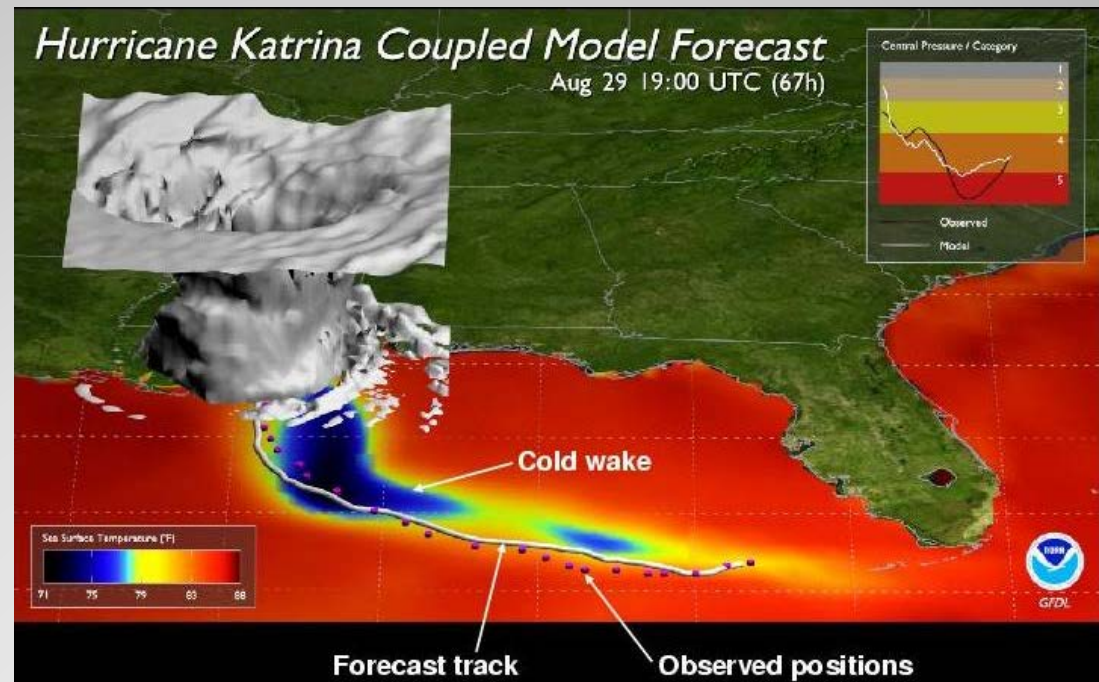
Sources of Vertical Mixing

- Wind-driven stirring
- Tidal generation of internal waves
- *Biota*
- *Hurricanes*



Sources of Vertical Mixing

- Wind-driven stirring
- Tidal generation of internal waves
- *Biota*
- *Hurricanes*



Sources of Vertical Mixing

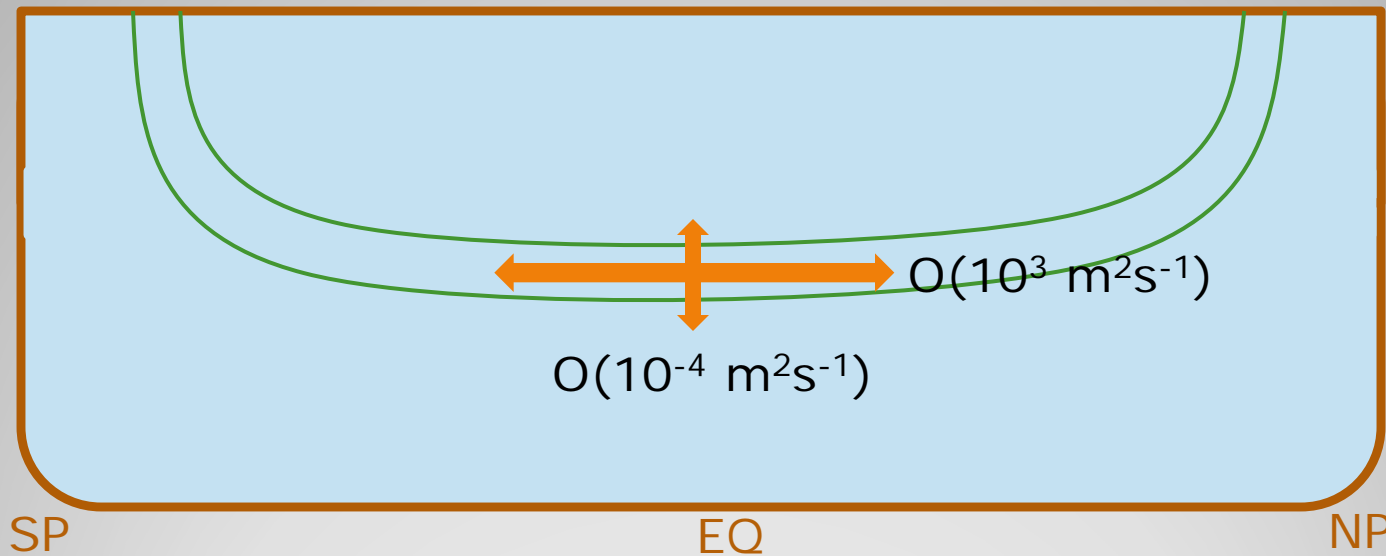
- Irrespective of the source can treat process as diffusion
 - Heat flux is proportional to temperature difference

$$\partial_z (K_v \partial_z T)$$

- K_v is a vertical eddy diffusivity

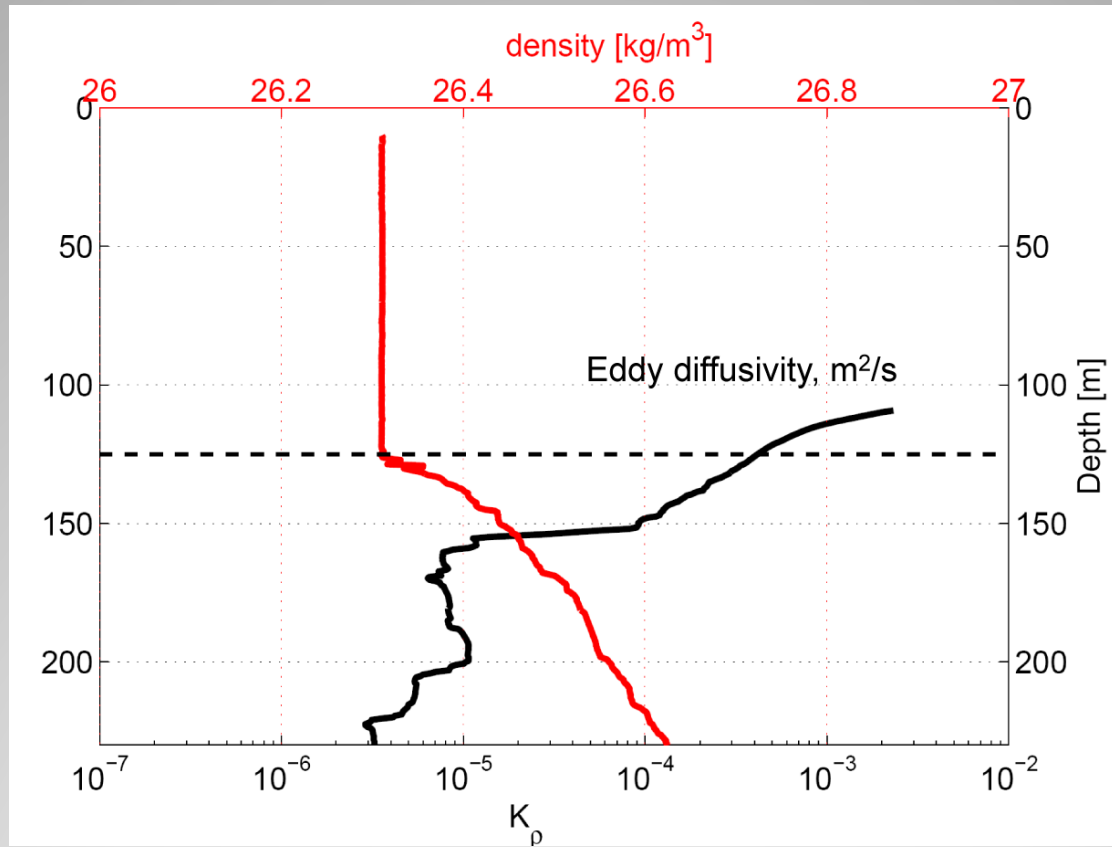
Eddy Diffusivity

- It is much easier to move along isopycnals than across them (diapycnal)



$$K_I/K_V = O(10^7)$$

Diapycnal vs. Isopycnal

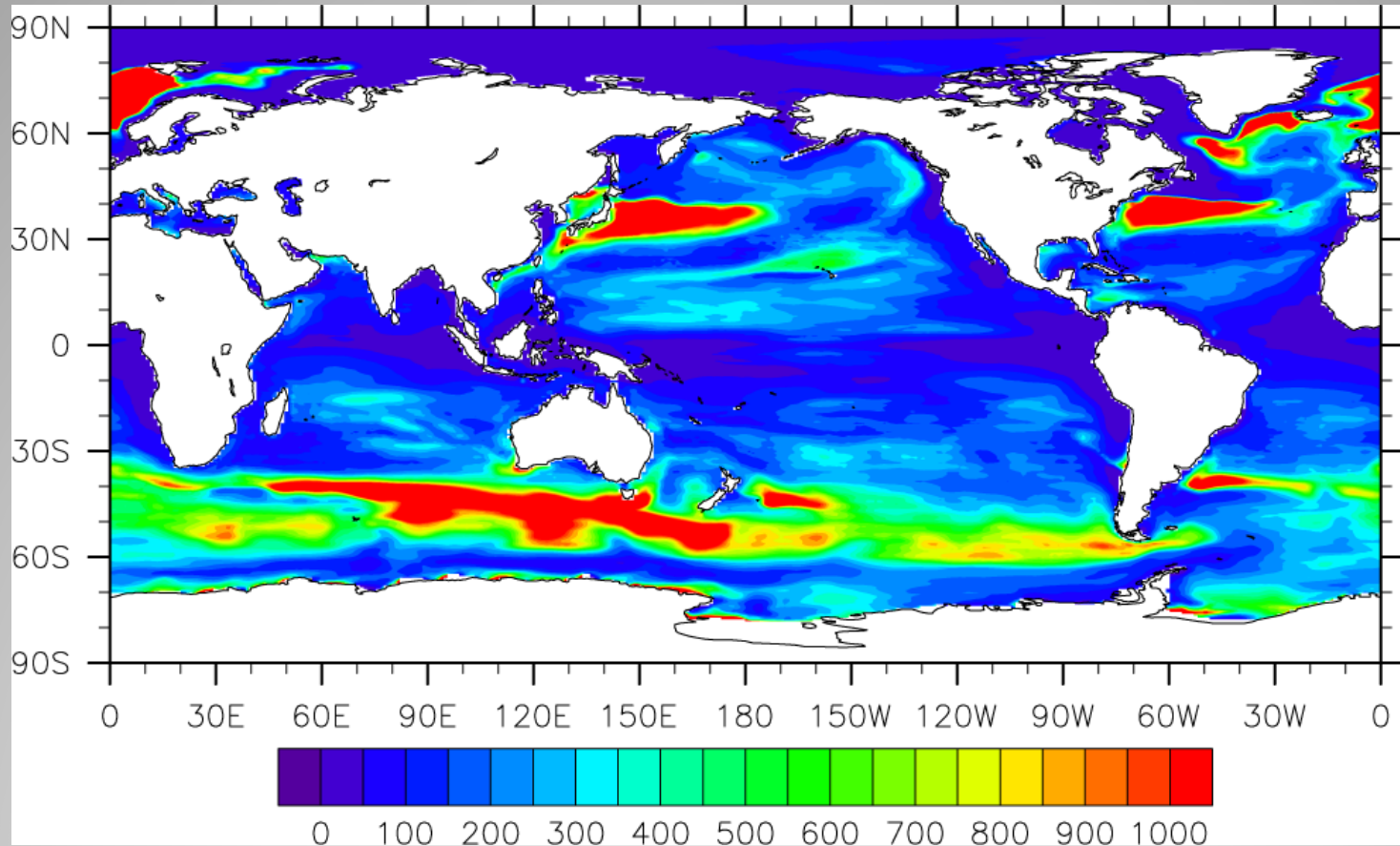


Results from the North Atlantic Tracer Release Experiment
Mixing changes with depth and stratification

Observations

- Philander & Pacanowksi, 1981 (K-Theory)
 - Vertical diffusion depends on Richardson Number, therefore stratification dependent
- KPP Scheme (Large et al., 1994)
 - Based off boundary layer principles learnt in meteorology
- Prescribed global background component
 - Increases with depth
 - Required for
 - numerical stability
 - fill gaps left by other parameterisations

Parameterisation of vertical mixing in a GCM



Average calculated K_v in top 100m ($10^{-4} \text{ m}^2 \text{ s}^{-1}$)

Vertical mixing field in a GCM

- The observations give a range of values of K_V for the ocean interior, but models use a single profile for the whole globe
- What impact would this uncertainty have on climate?

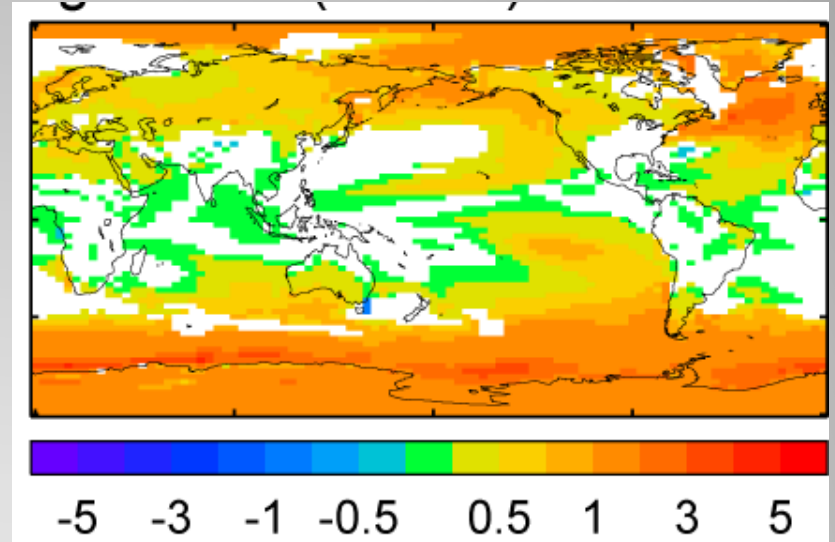
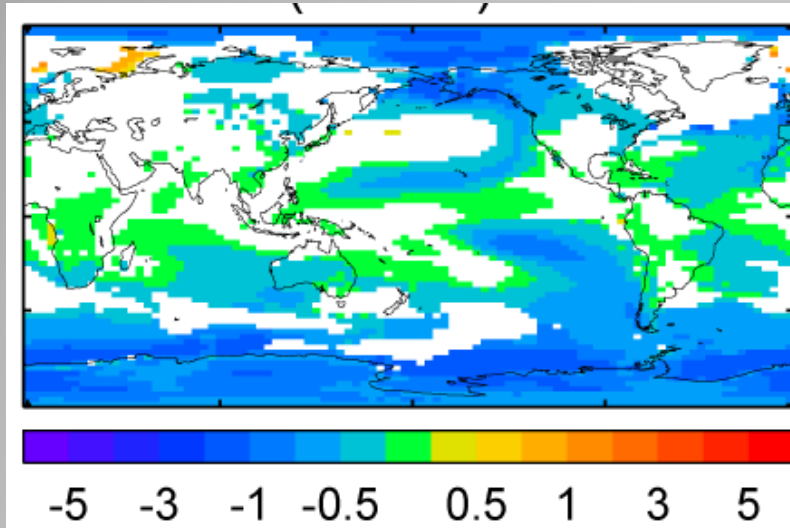
Climate Model Experiments

K_V (m^2s^{-1})	Low K_V Run	Standard K_V	High K_V Run
Surface	0.5×10^{-5}	1×10^{-5}	2×10^{-5}
Bottom	4×10^{-5}	15×10^{-5}	50×10^{-5}

Test of Uncertainty

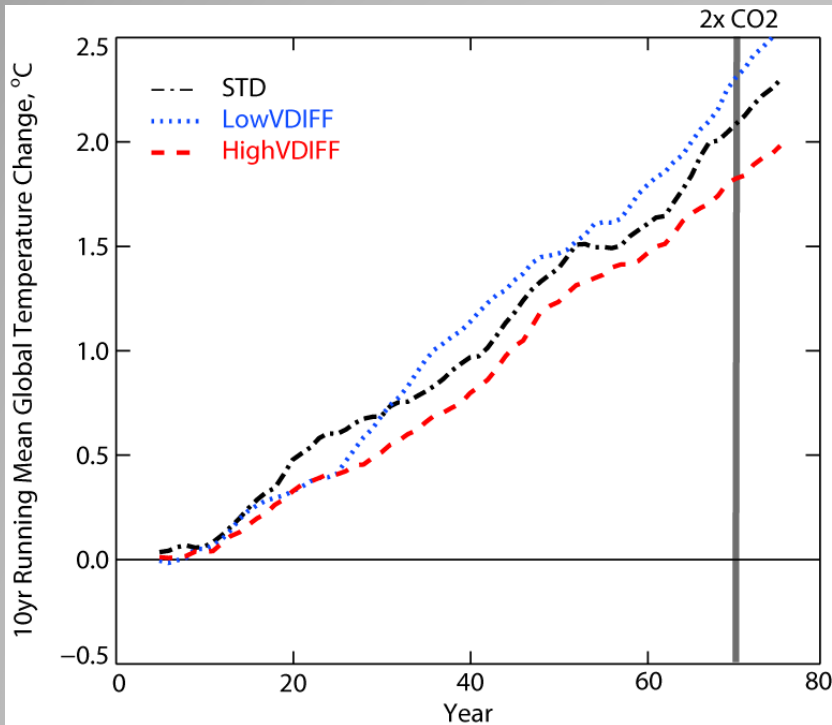
Low K_v

High K_v



Higher mixing leads to stronger ocean circulation
(c.f. with Sandström thoughts from earlier)

Impact on Mean Climate (0.8°C)



Run	Transient Climate Response (°C)
Low K_v	2.3
Standard	2.1
High K_v	1.8

- Transient response opposite from equilibrium one
- Higher mixing means better draw down of heat from the ocean surface

Impact on Climate Projections

- Uncertainty in K_v leads to significant uncertainty in climate models and their projections of climate change
- How can we constrain that uncertainty?
 - Look at individual sources of mixing
 - Determine their impact
 - Observe the sources directly rather than attempt to measure uncertainty
 - *Think about using paleoclimates*

Making a more certain world

- What are the sources of mixing in the ocean?
 - Wind-driven stirring
 - *Relatively well constrained*
 - *Already parameterised in climate models*
 - Tidal generation of internal waves
 - Biota
 - Hurricanes

Sources of Ocean Mixing

Tropical Cyclones
(hurricanes,
typhoons etc)

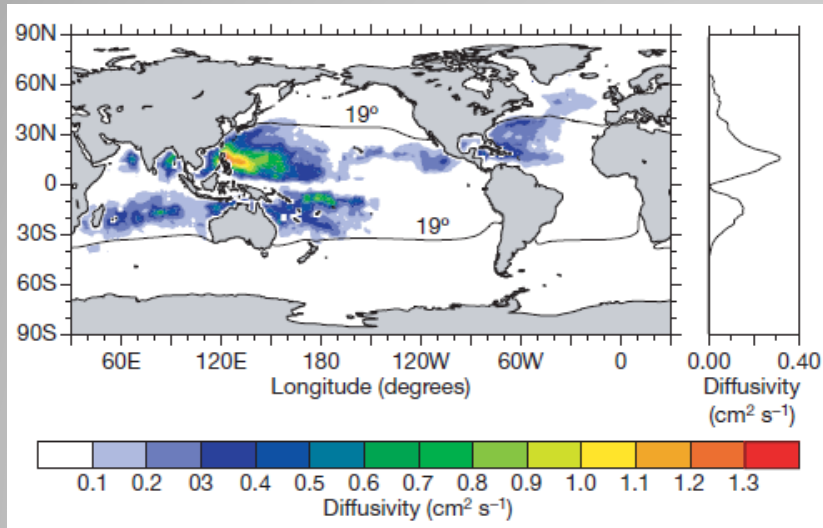
Impact on ocean
circulation is
debated, as they
are short, transient
features

Hurricanes

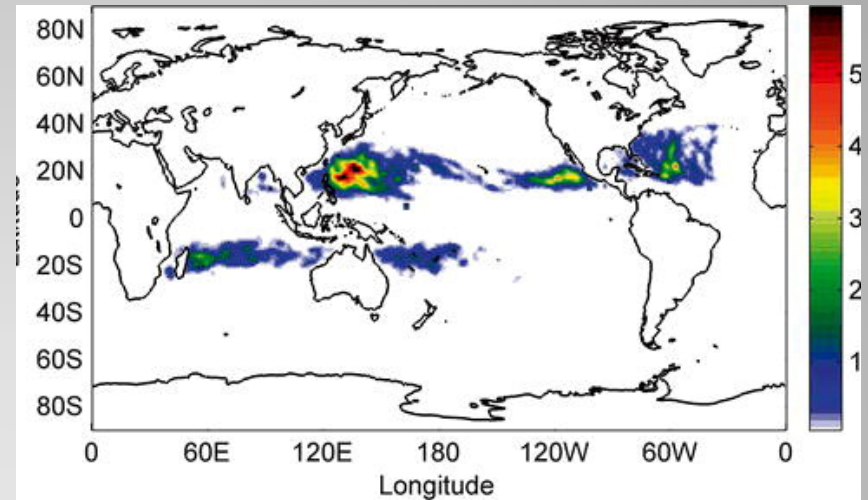


Striver & Huber, 2007

Liu et al, 2008

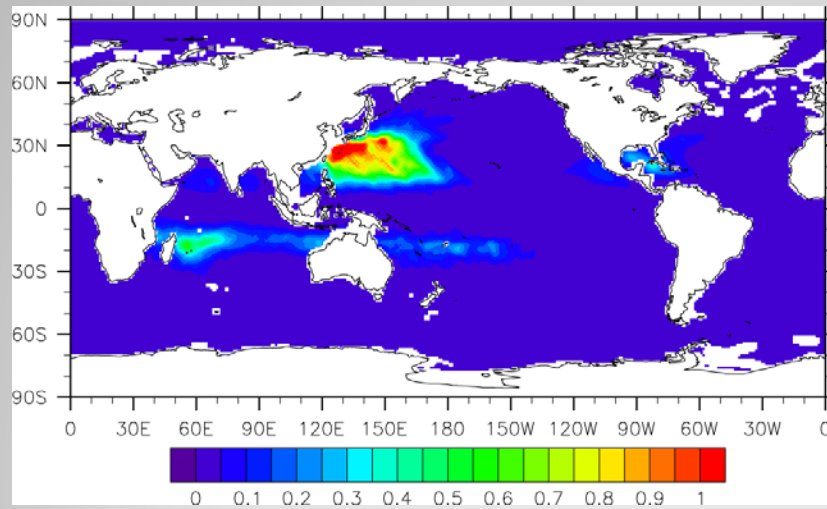
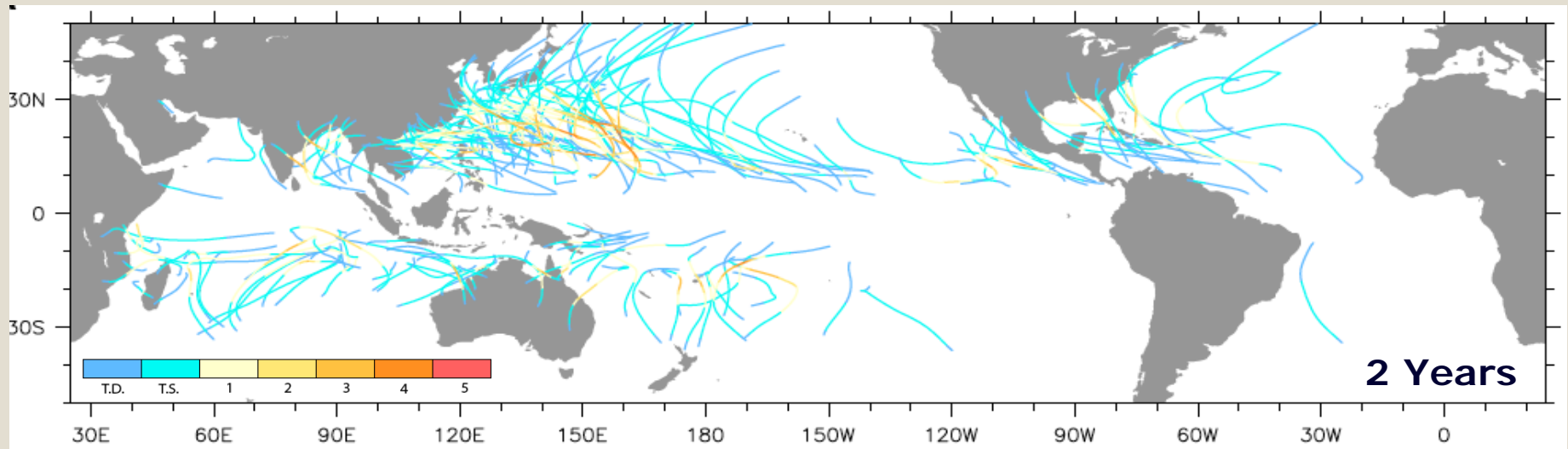


Maximum of $10^{-4} \text{m}^2 \text{s}^{-1}$



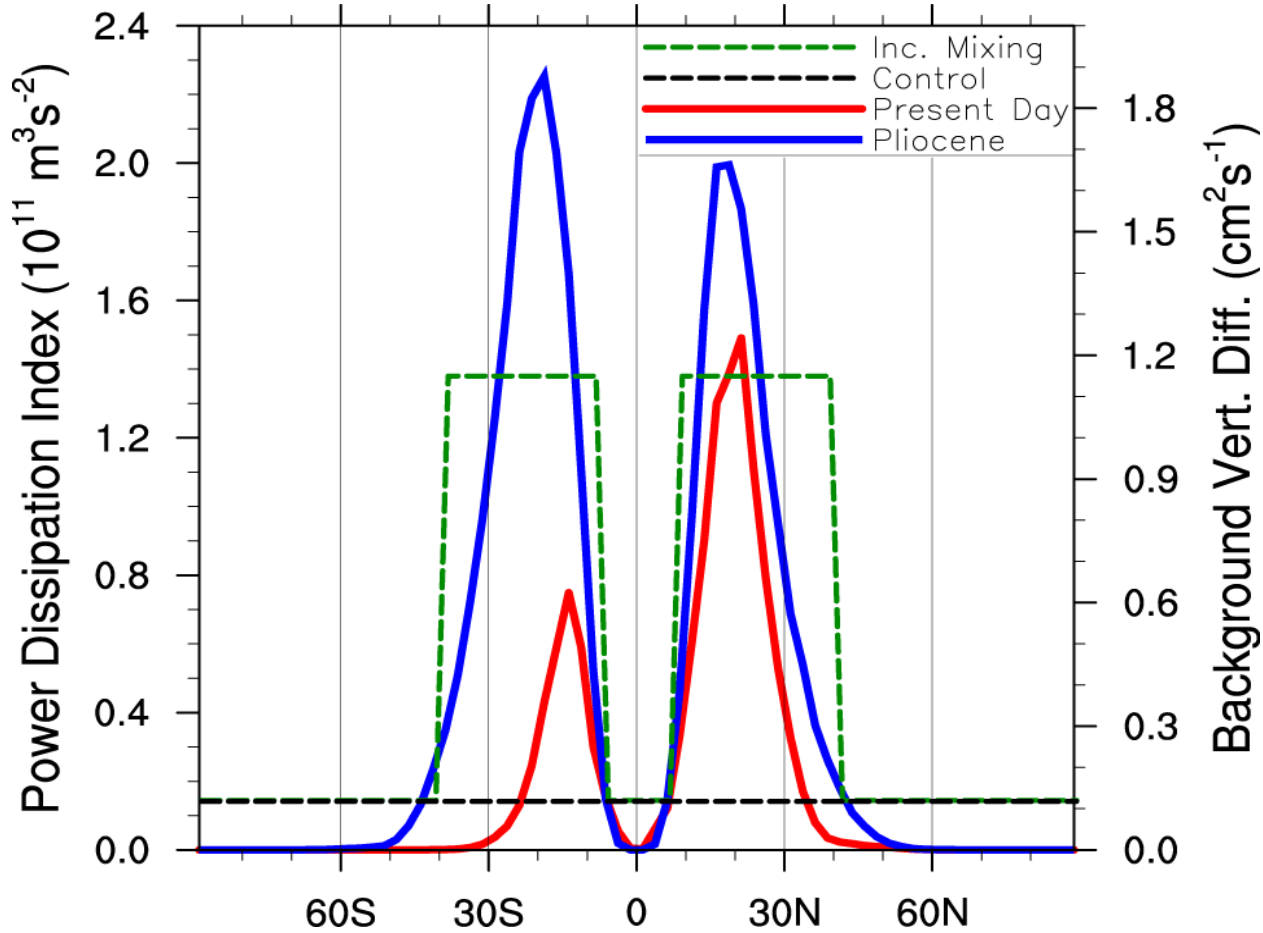
Maximum of $5 \times 10^{-4} \text{m}^2 \text{s}^{-1}$

Hurricane K_v Fields - Observed



Calculate diffusivity
 from energy balance of
 mixed layer (Georgy)
Maximum of $10^{-4}m^2s^{-1}$

Reconstructed Hurricanes



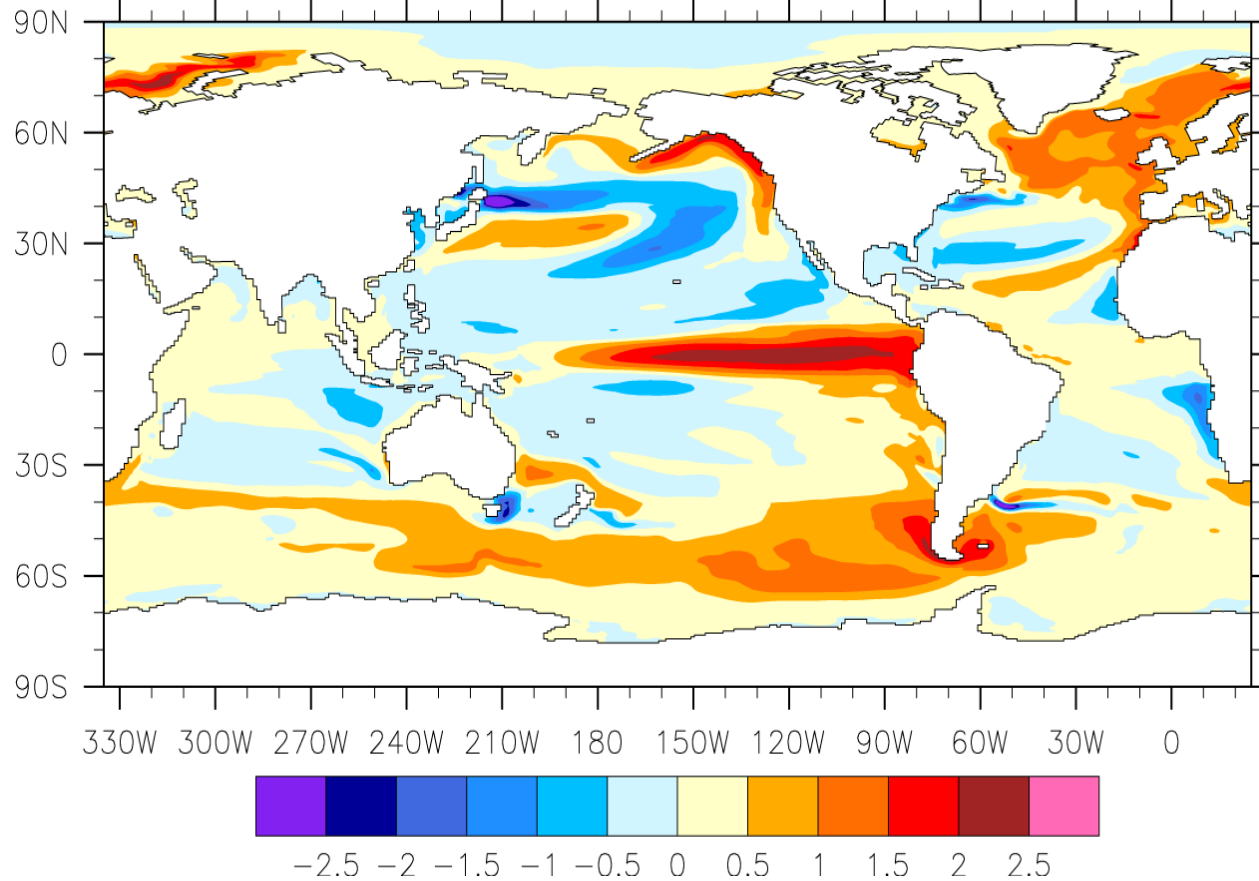
Include hurricanes as two stripes of extra mixing in their common latitudes

Use 10x normal background mixing to test sensitivity

Including hurricanes in a GCM

TC (200m, 8-40N/S) - Control

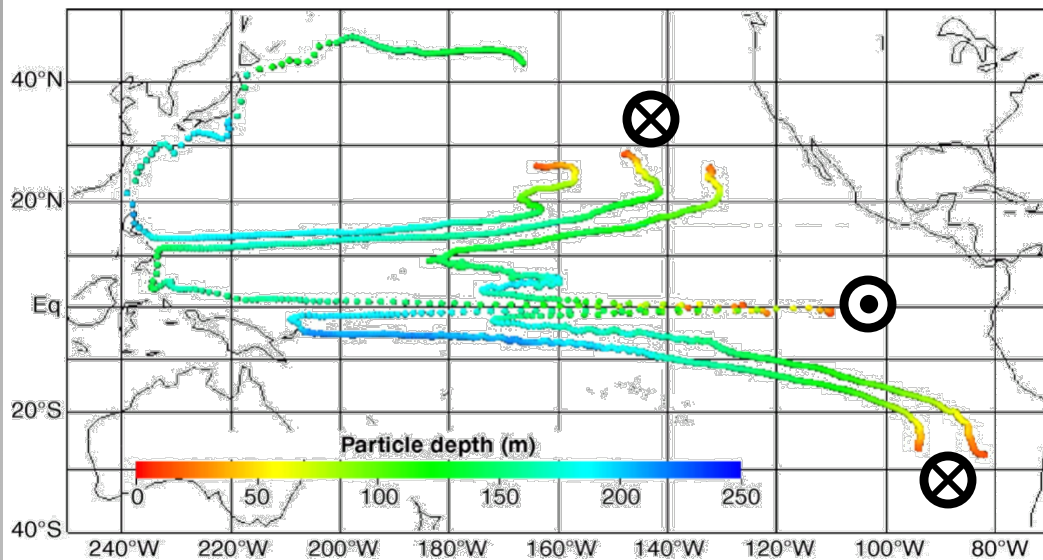
°C



Inclusion of hurricanes in coupled climate model causes up to 2.5°C

El Niño region is most strongly affected

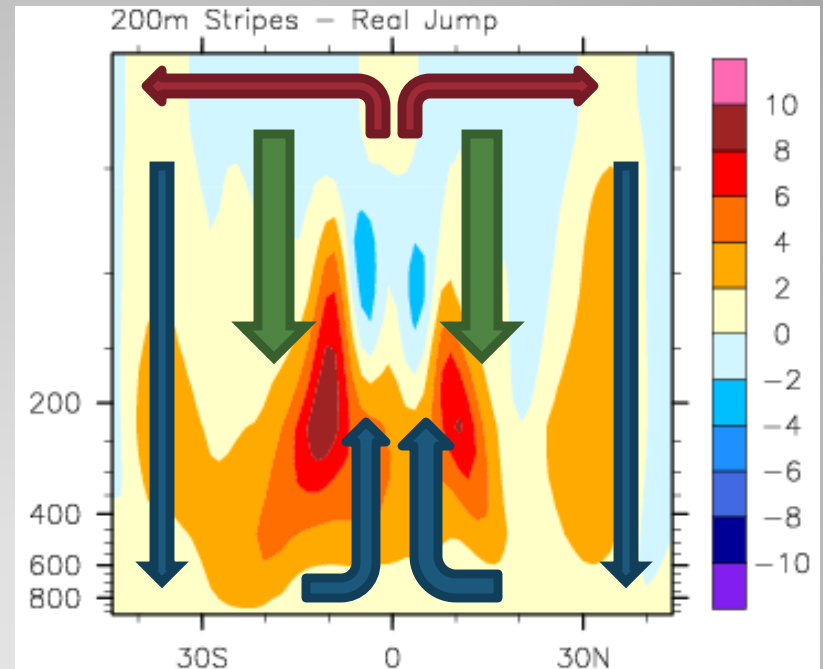
Significant surface warming



- Water is subducted in subtropical East Pacific
- Travels west towards warm pool
- Catches EUC and upwells in cold tongue
- From Gu & Philander '97

The Wind-driven Subtropical Cell

- Hurricanes act to mix warm water from surface to ~200m
- Warms the return flow of the cell



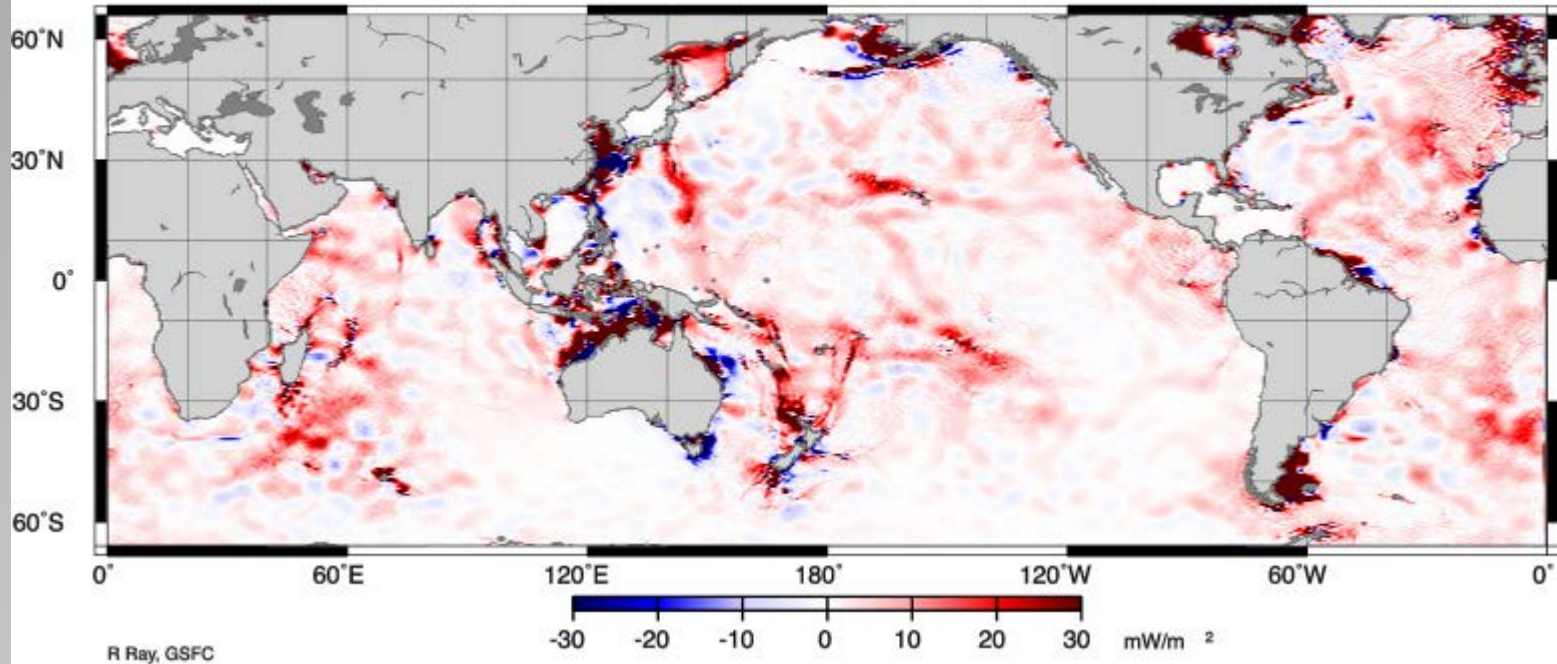
Cross section at 180°
(before EUC hits the surface)

Mechanism of hurricane warmth

- What are the sources of mixing in the ocean?
 - Wind-driven stirring
 - *Relatively well constrained*
 - *Already parameterised in climate models*
 - Tidal generation of internal waves
 - Biota
 - Hurricanes
 - *Could be important*
 - *Needs a lot of work – get a grad student to do it*

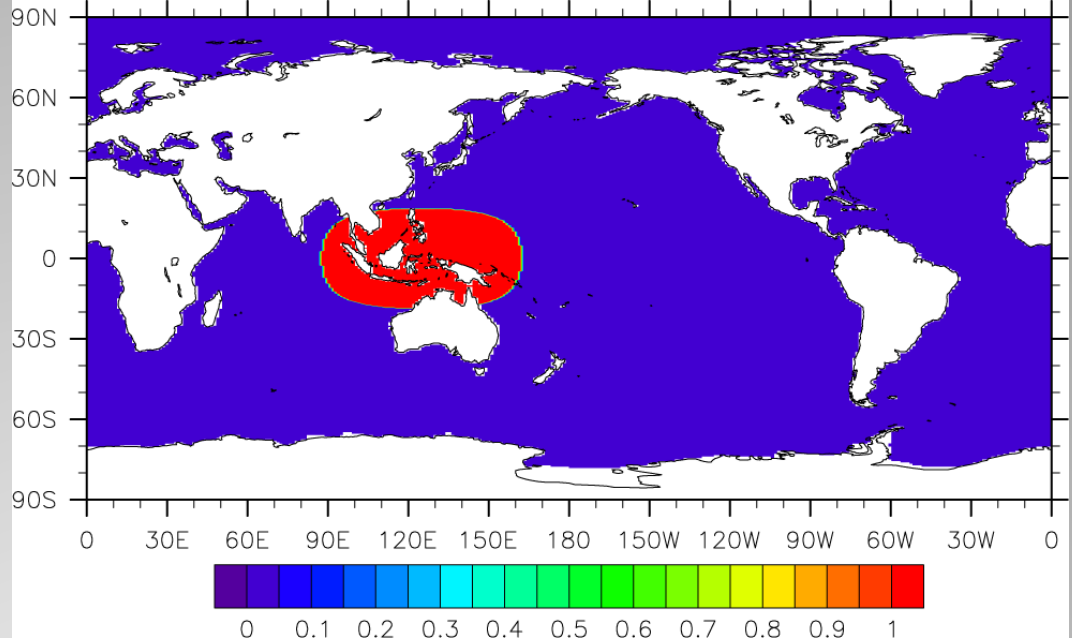
Sources of Ocean Mixing

M2 Tidal Energy Dissipation
From balance of working and flux divergence



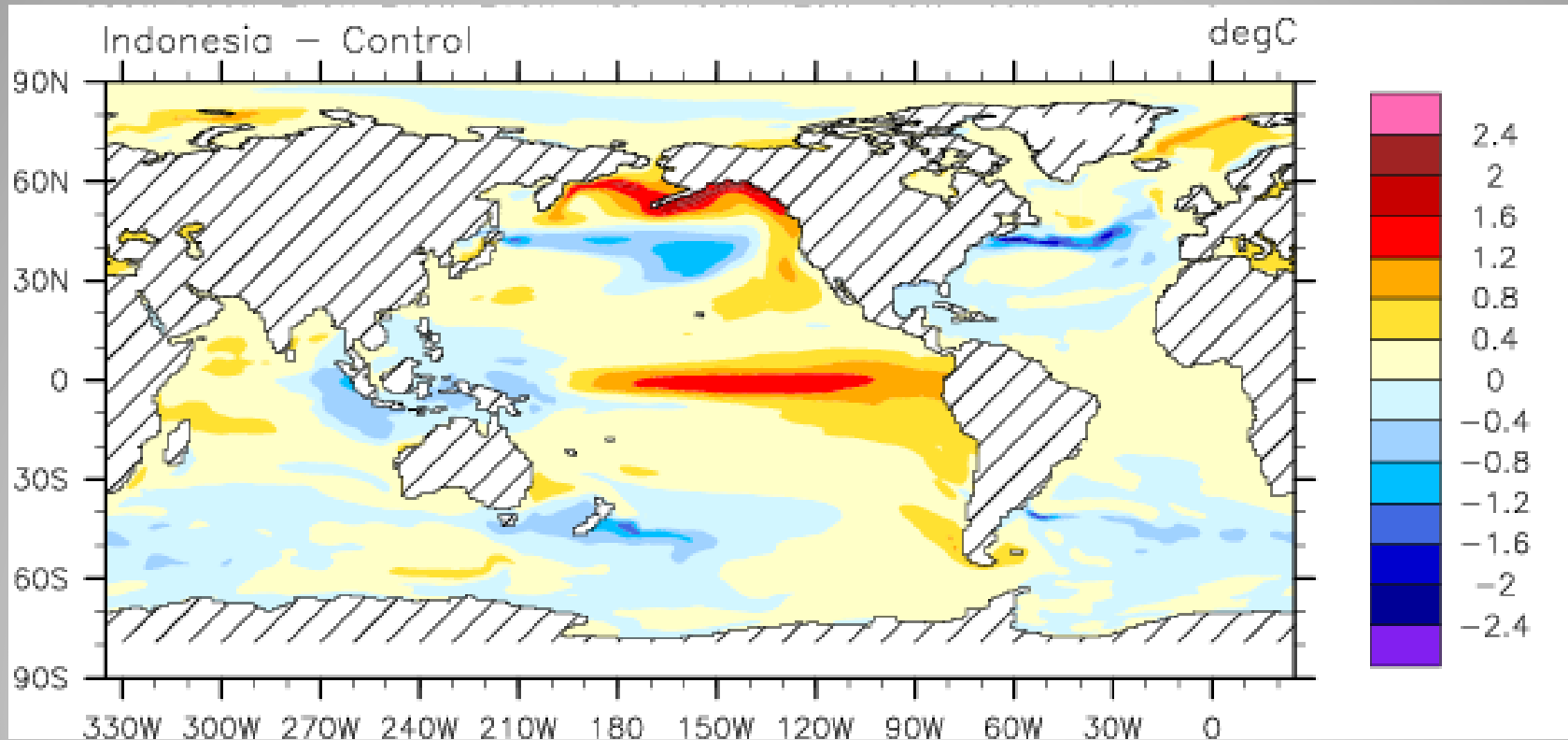
Spatial pattern of tidal energy

- Add area of increased mixing over Indonesia from surface to 200m
- 1/1000th of maximum calculated value



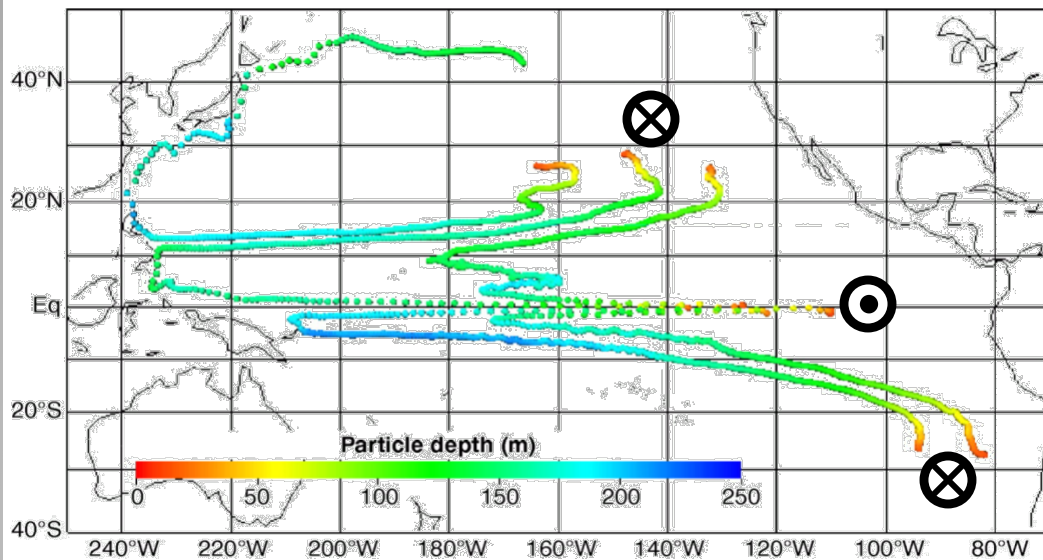
Additional imposed K_v averaged
over top 100m ($10^{-4} \text{ m}^2\text{s}^{-1}$)

**Is changing bathymetry
important?**



SST maximum in Tropics $\sim 1.5^{\circ}\text{C}$ (less than hurricanes)
Impacts in North Pacific, but run is not equilibrated yet

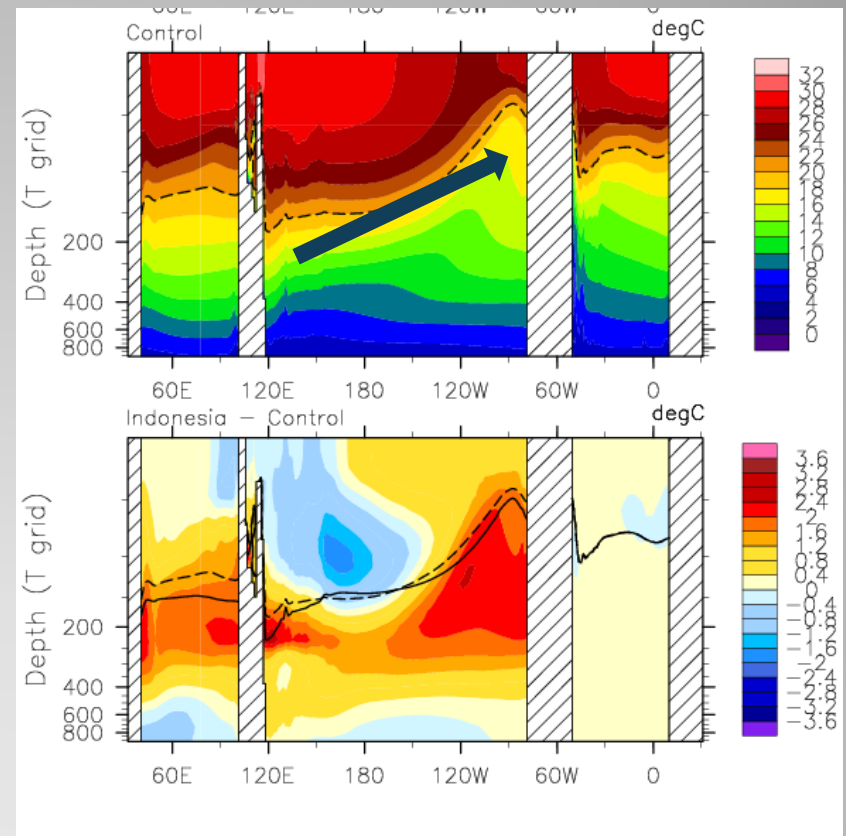
Impact on SST



- Water is subducted in subtropical East Pacific
- Travels west towards warm pool
- Catches EUC and upwells in cold tongue
- From Gu & Philander '97

The Wind-driven Subtropical Cell

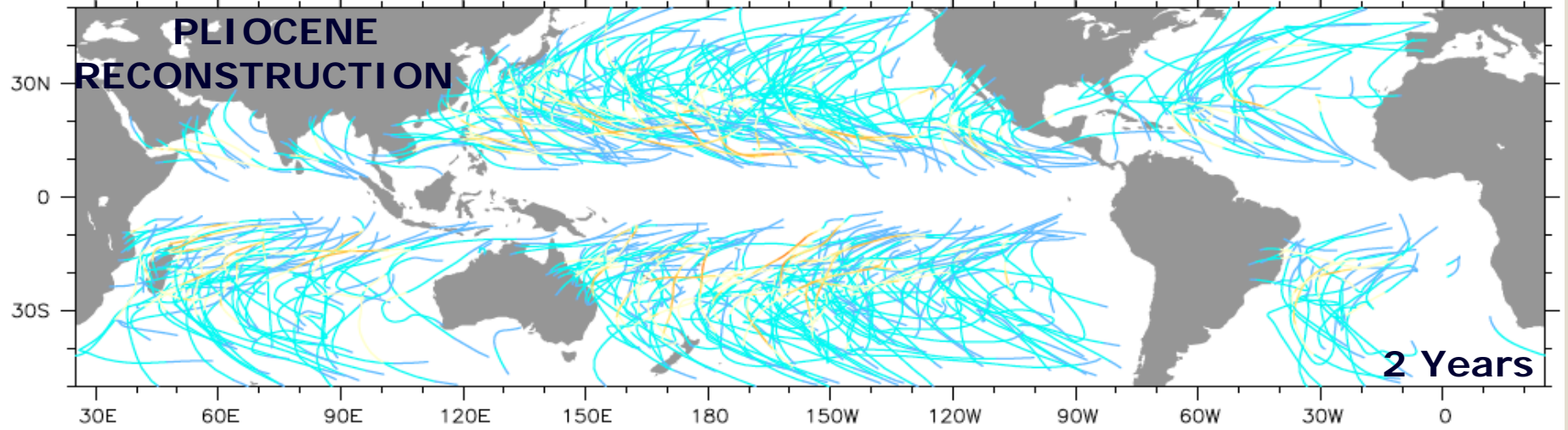
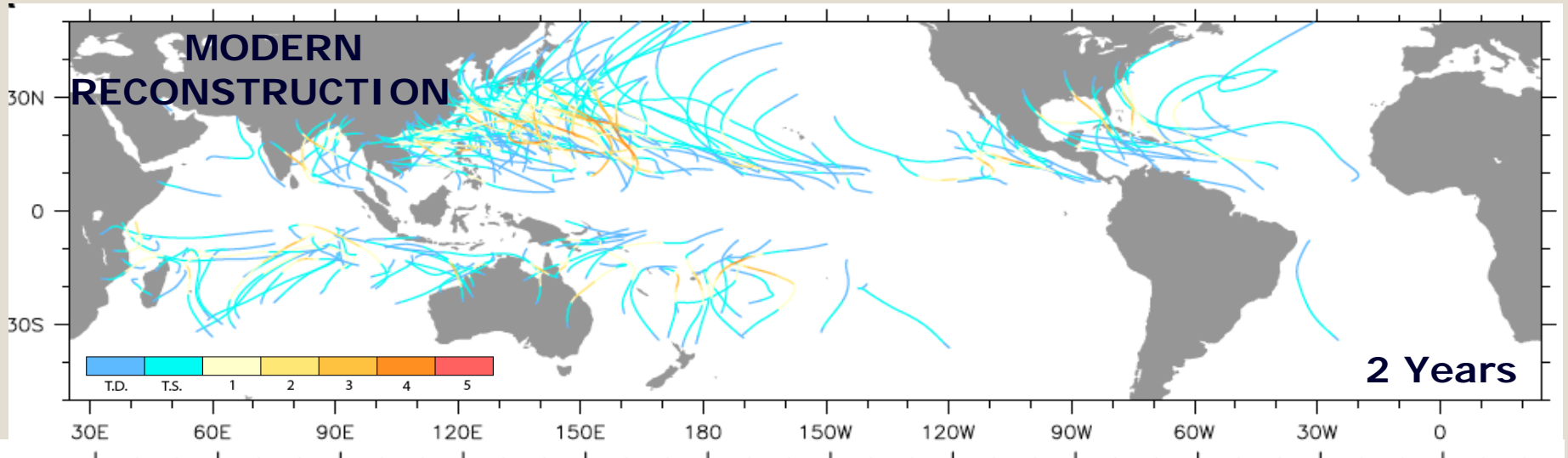
- Tidal mixing over Indonesia deepens mixed layer
- More warmth enters EUC
- Warmer upwelling
- Only impacts top of EUC, which reaches surface further west



Mechanism of Tidal Warmth

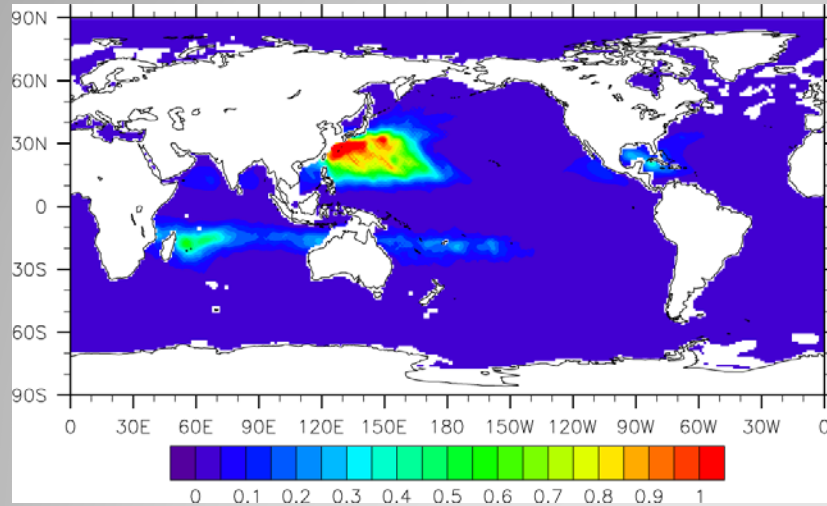
- There is vertical mixing in the ocean, which comes from a variety of sources
- Uncertainty & changes in vertical mixing in the ocean has climatic consequences
- Considering the sources individually may reduce that uncertainty (rather than having a single term do it all)
- Ocean could be sensitive to both tides and hurricanes

Conclusions



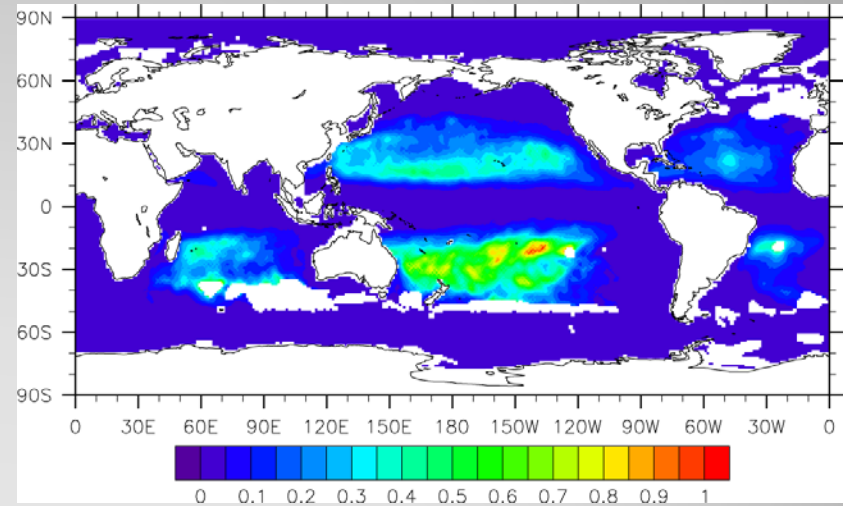
Reconstructed Hurricanes

Present Day



Maximum of $10^{-4}\text{m}^2\text{s}^{-1}$

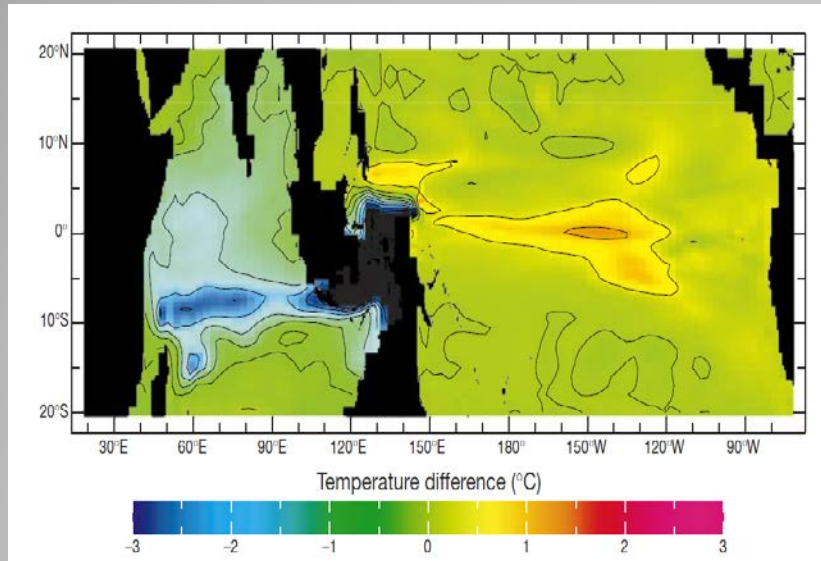
Pliocene



Maximum of $10^{-4}\text{m}^2\text{s}^{-1}$,
much broader bands

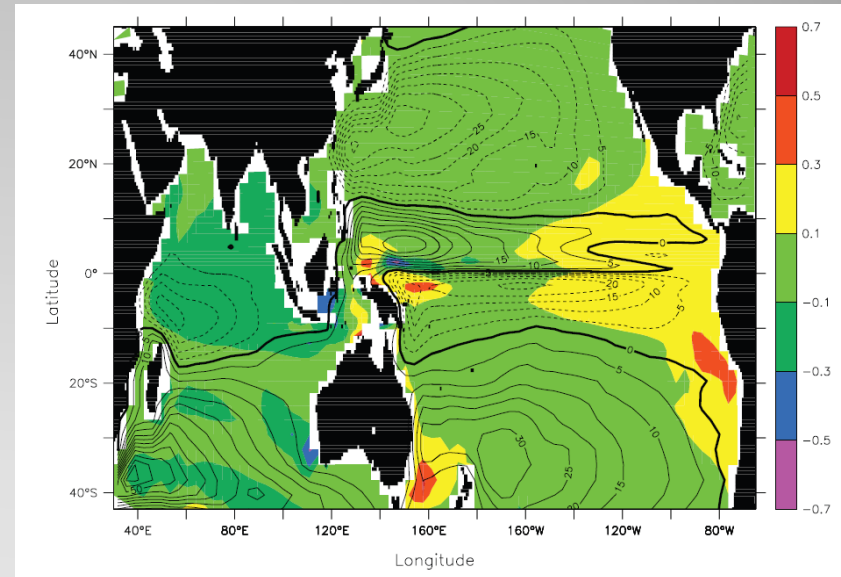
Hurricane K_v Fields - Modeled

Cane & Molnar, 2001



Maximum of 1°C at 100m

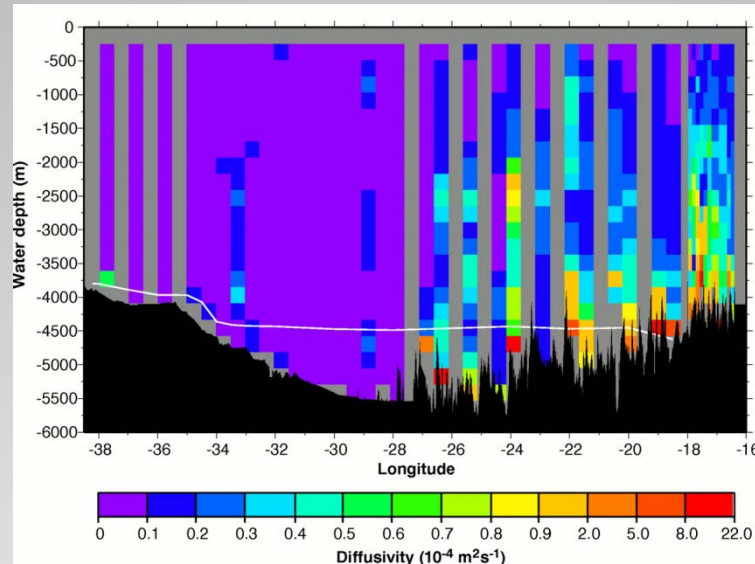
Jochum et al, 2009



Maximum of 0.1°C at 150m

Climate simulations with altered Indonesian bathymetry

- Moon's Orbit is stable
- Roughness of the ocean bottom is not:



- Changes in Isthmus of Panama and Indonesia over past 5 Myrs

Changes in Tidal Mixing?